

The representation of hydrological dynamical systems using Extended Petri Nets (EPN)

Marialaura Bancheri, ¹Francesco Serafin² and Riccardo Rigon²

¹Institute for Mediterranean Agricultural and Forestry systems (ISAFOM), National Research Council (CNR), Ercolano (NA), Italy ²Department of Civil, Environmental and Machanical Engineering, University of Tranto, Italy

 2 Department of Civil, Environmental and Mechanical Engineering, University of Trento, Tento, Italy

Key Points:

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- We present a graphical system to represent hydrological dynamical systems called Extended Petri Nets (EPN).
- EPN have a one-to-one correspondence with the equations that drive systems.
- EPN topology and connections clarify the causal relationship between compartments and the feedback between them. Two different types of feedback are presented.
 - EPN can be used to formalize perceptual models from field work into equations.

Corresponding author: Riccardo Rigon, riccardo.rigon@unitn.it

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15 Abstract

This work presents a new graphical system to represent hydrological dynamical mod-16 els and their interactions. We propose an extended version of the Petri Nets mathemat-17 ical modelling language, the Extended Petri Nets (EPN), which allows for an immedi-18 ate translation from the graphics of the model to its mathematical representation in a 19 clear way. We introduce the principal objects of the EPN representation (*i.e.* places, tran-20 sitions, arcs, controllers and splitters) and their use in hydrological systems. We show 21 how to cast hydrological models in EPN and how to complete their mathematical de-22 scription using a dictionary for the symbols and an expression table for the flux equa-23 tions. Thanks to the compositional property of EPN, we show how it is possible to rep-24 resent either a single hydrological response unit or a complex catchment where multi-25 ple systems of equations are solved simultaneously. Finally, EPN can be used to describe 26 complex earth system models that include feedback between the water, energy and car-27 bon budgets. The representation of hydrological dynamical systems with EPN provides 28 a clear visualization of the relations and feedback between subsystems, which can be stud-29 ied with techniques introduced in non-linear systems theory and control theory. 30

31 1 Introduction

In the broad array of hydrological models (Beven, 2011; Wagener, Wheater, & Gupta, 32 2004) an important category comprises those models that solve systems of Ordinary Dif-33 ferential Equations (ODEs) and their discrete counterparts (for an overview, please re-34 fer to Singh and Woolhiser (2002) and Kampf and Burges (2007)). This category includes 35 lumped models, that is to say, models where spatial hydrological variability is integrated 36 over single elements called Hydrological Response Units (HRUs): each HRU represents 37 a certain sub-catchment, while the spatial organization of basins, if required at coarser 38 scales, is obtained by connecting HRUs as nodes of a network. In this case, lumped mod-39 els are also called "integral distributed models", (Todini, 1988). In each HRU, a model 40 can treat the internal processes (runoff, evapotranspiration, root zone moisture, and so 41 on) by using one or more ODEs. Therefore, integral distributed models are formed by 42 systems of systems of ODEs. 43

Not all the aforementioned elements are present in all hydrological models, nor is 44 the same nomenclature used. However, if we take as an example the models collected 45 in the MARRMot 1.0 toolbox (Knoben, Freer, Fowler, Peel, & Woods, 2019), we have 46 a substantial group (46) of the most widely used hydrological models, all of which solve 47 ODEs. In literature, (Birkel, Soulsby, & Tetzlaff, 2011; Fenicia, Savenije, Matgen, & Pfis-48 ter, 2008; Hrachowitz, Savenije, Bogaard, Soulsby, & Tetzlaff, 2013), these Hydrologi-49 cal Dynamical Systems (henceforth HDSys) are used to interpret any of the hydrolog-50 ical processes from hillslope to catchment scale: they are ubiquitous. 51

The great variety of available models draws attention to the need to find some mathematical criterion for diagnosing their differences (e.g., Clark et al. (2008)). In this paper we suggest that associating an appropriate graphical-mathematical representation to each model can be a part of the diagnostic process.

Graphical representation has been fruitful in the sciences: the epitome is the case 56 of Feynman diagrams in quantum electrodynamics (Kaiser, 2005), but representations 57 of electrical circuits (Lohn & Colombano, 1999), stock-flow diagrams of system dynam-58 ics models are also good examples (Takahashi, 2005) and reaction networks (Baez & Pol-59 lard, 2017; Herajy & Heiner, 2015) are also interesting examples. The resulting theories, 60 informed by the diagrams, differed significantly from earlier approaches in the way the 61 relevant phenomena were conceptualized and modelled. We believe that devising a graph-62 ical representation for hydrological models can also be fruitful, especially if the graph-63 ics are more than pictorial representations. As Oster, Perelson, and Katchalsky (1971) 64 suggest, we seek a system where the dynamical equations can be read algorithmically 65

from the graphs and diagrams, which are actually another notation for the equations themselves.

In hydrology we have great demands as we deal with various dynamical systems besides the water budget, such as the energy budget, the travel time transport of water, and the carbon cycle, to name a few. Therefore, the graphical representation developed should be expandable to more than one of the Earth system cycles; it should imply their mathematics; and it should help visualize their reciprocal feedbacks.

In our work, we want to complement the work presented, for instance, in Fenicia
and Kavetski (2011) and in Clark et al. (2015). Those are papers with a large scope, and
they treat very broad questions, from how to infer a model's structure using heuristic
analyses of the functioning catchment (e.g., Butts, Payne, Kristensen, and Madsen (2004))
to the numerics used in sound, high-performance tools. With respect to the models addressed by those papers, the approach of this paper is agnostic: it does not explain how
to build models but aims to present them in a clear way.

In summary, our paper tries to answer the following questions: is there a good way 80 to graphically represent budgets (water, energy and other) that gives a clear idea of the 81 type of interactions they are subject to before seeing the equations? Where in a graph-82 ical representation can information about fluxes and parameters be optimally placed? 83 Can we obtain a graphic language that corresponds to mathematics in a strict and uni-84 vocal manner? Can the graphical representation help translate the perceptual models 85 derived from field work into mathematics and equations? Can we visually represent the 86 feedbacks between hydrology and ecosystems? 87

2 Examples of graphical representation of hydrological models

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To expound what was said in the Introduction, we reproduce here figures representing some well known hydrological models.



Figure 1. Representation of the model proposed in Birkel et al., 2011. The figure is adapted from Soulsby et al. (2016)

Figure 1 shows a schematic representation of the model proposed in Birkel et al.

(2011), which we shall refer to as the BST model (after Birkel, Soulsby, Tetzlaff). In the

graphic, the relationships between different BST parts are clear; this is not true for the

⁹⁶ fluxes, which have their mathematical expressions annotated in the graphic. Computer





scientists would say that the figure has been given too many responsibilities and too much
information, resulting in a cluttered graphic. To understand and reproduce the BST model,
decryption work is required in a back and forth process between the image and the text.
This process is probably unavoidable in all cases, but the reading can be made easier by
referring to standard places in the manuscript.

Figure 2 refers to the Hydrologiska Byråns Vattenbalansavdelning (HBV), adapted 104 from Seibert and Vis (2012), a standard reference for HBV. Those Authors opted for a 105 pictorial representation that cannot be considered very explicative from a mathemati-106 cal point of view, as it serves to identify the compartments of Earth surface involved. 107 While the Figure is very effective in providing an immediate association between the model 108 components and their natural counterparts, the interested reader must, however, peruse 109 other papers to get all the information needed to understand the workings of the HBV 110 model. 111

Figure 3, adapted from Hrachowitz et al. (2013), is one of three model structures used in a heuristic procedure (Fenicia et al., 2008) to assess catchment behaviors. The figure conveys a lot, but details about flux partition remain unclear. Single reservoirs need to act like two or three reservoirs, as represented by the use of different colours. The (inattentive) reader could be easily confounded to see only four reservoirs in this model, when, instead, the S_U reservoir should be split in two, and some others are missing too, as we shall see later.

The model representations in Figures (1) to (3) keep some elements fixed, namely, the reservoirs and the arrows. Others elements vary, and some are discarded, in accordance with the Authors' views. That is to say, it is not possible to gather the main information at a glance or, rather, there is no common understanding of what the main information to be conveyed is. We cannot easily see the similarities between models, and the style changes in representation make any understanding even more difficult.



Figure 3. Representation of one of the models proposed in Hrachowitz et al. (2013), here called the Ard-Burn model

The goal of this paper is to bring order to HDSys representations by building an algebra of graphical objects where any symbol will correspond to a mathematical term or group of terms. The main information to be communicated is the number of equations that a model uses and the number of input and output fluxes present for each equation. At the same time, the number and location of model parameters should be clear, but, in our opinion, need not be communicated directly by the graphics.

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3 Principal graphical objects in Extended Petri Networks

Among the various possible graphic representation, we find that the Petri Nets (PN) 134 are particularly suited to our scope. PN are a mathematical modelling language for the 135 description of distributed systems. The concept was originally presented in Carl Adam 136 Petri's dissertation (Petri, 1966) and their early development and applications are found 137 in reports that date back to the 1970s. PN became popular in theoretical computer sci-138 ence (Jensen & Kristensen, 2009), biology (Koch, 2010; Koch, Reisig, & Schreiber, 2010; 139 Wilkinson, 2011), especially to represent parallel or concurrent activities (Murata, 1989), 140 stochastic mechanics (Baez & Biamonte, 2012; Haas, 2006; Marsan, Balbo, Conte, Do-141 natelli, & Franceschinis, 1994) and to describe reaction networks (Gilbert & Heiner, 2006; 142 Herajy & Heiner, 2015). In the case of reaction networks, clearly treated in Herajy and 143 Heiner (2015), there are specific rules for computation, which are implicit in the PN struc-144 ture used, that do not lead to correct mass and energy budget equations. This matter 145 is referred to in more detail in the supplementary material of this paper. 146

Initially, PN were used to model discrete time processes managing discrete, numerable quantities. However, HDSys require a time-dependent form of PN. Such a form is already present in literature, (Alla & David, 1998; Berthomieu & Diaz, 1991; Champagnat, Esteban, Pingaud, & Valette, 1998; Merlin & Farber, 1976; Ramchandani, 1974) and is usually called "Time Continuous Petri Nets". These are the generalization of discrete processes that are approximated as continuous ones (Silva & Recalde, 2004). However in HDSys, we mostly deal with systems of ODEs, where the equations are usually non-







linear and the state variables are inherently continuous (mass, energy and momentum
of water or other substances). Thus we required a different type of PN that we have
called Extended Petri Nets (EPN), with different rules from, for example, the reaction
networks or other typologies of PN.

When looking at PN, hydrologists must adjust their interpretative habits: reser-158 voirs (now called **places**) are represented as circles, and fluxes (now called **transitions**) 159 between reservoirs (or places) are represented as squares. To distinguish between differ-160 ent places, the graphical objects can be colored; conventionally, we use the same color 161 for places and transitions describing the same compartment, such as, for instance, the 162 soil or root zone as distinct from the groundwater zone. The graphical objects have enough 163 space for the symbol of the variable they deal with, as shown in Figure 4. A third group 164 of objects are the **controllers** (represented by a triangle). They are quantities that af-165 fect fluxes but are not fluxes themselves. Their value can depend on one or more state 166 variables, i.e. on places, and they are in charge of regulating fluxes. As an example of 167 a controller, consider a mass flux, Q, proportional to the storage, S, such that Q = kS. 168 If k = k(T), where T is the temperature, then T is a controller of the flux. 169

The connection between places and transitions is shown with an **arc**; arcs between 171 two places (reservoirs) or between two transitions (fluxes) are not allowed. As shown in 172 Figure 4, arcs can be drawn in different ways to convey more detail: if they carry a lin-173 ear flux they are generic and do not include any symbols; if the carry a non-linear flux, 174 they are marked by a coloured bullet. Binding arcs are used when two different fluxes 175 in two different budgets contain the same variable. That is to say, they join two tran-176 sitions that contain the same variable for graphical reasons, such as, for example, evap-177 otranspiration in the water and energy budgets, as shown in section 7. Oriented dashed 178 arcs show connections from places to controllers and from controllers to transitions. Con-179 nections between places and transitions that pass trough controllers only affect the ex-180 pression of fluxes but do not alter the number of equations. Any oriented arc also rep-181 resents a causal relation between the originating entity and the receiving one: upstream 182 quantities can be thought to cause downstream ones. Therefore the controllers show the 183 causal relationship between state variables and fluxes, which would otherwise be hidden 184 graphically. For this reason we call the wiring from places to controllers to transitions 185 hidden wiring or **h-wiring**, while the wiring that connects directly between places and 186 transitions is called flux wiring or **f-wiring**. 187

In Figure 4 we also introduce a small, solid, black circle, which is used to mark a measured quantity, i.e. a quantity that is given as known input and drives the simula-

tion. The most common example of known input is precipitation, which is usually ob-190 tained from ground measurements or other sources. The small, empty circle represents 191 a quantity that is also given but is used to assess the goodness of the model. In hydrol-192 ogy, the typical case is the discharge, which is an output of the models and whose mea-193 sured values are used for validation. The big circle with the dotted border represents in-194 stead a hidden place whose budget is stationary, as it returns all the mass it takes in. 195 A typical example in hydrological models is that of uphill surface waters and ground-196 waters summing to give the total surface discharge. 197

198 All the allowable connections between EPN objects are represented in Figure 5; no other type of connection is possible. A transition can be connected to more than one place, 199 implying the existence of a partition coefficient, represented by a **splitter** (the diamond 200 symbol in Figure 4). For instance, the total amount of precipitation can be divided into 201 snowfall and rainfall, or between two reservoirs representing surface waters and the root 202 zone. In those cases the splitter represents the need for some rule to separate the fluxes. 203 Figure 5 shows a splitter in action, where precipitation J is divided into 2 components, 204 J_{up} and J_{down} . In the case presented in section 4.1, the separation is simply obtained 205 with a partition coefficient, for which α part of the precipitation goes into a surface reser-206 voir and $(1-\alpha)$ part goes to a soil reservoir. Usually, however, each internal transition 207 is connected to only one place. Similarly, a place can be connected to more than one tran-208 sition, also implying a partitioning rule or coefficient. Two places cannot be connected to a unique transition, and this marks a substantial difference with reaction networks 210 (Gilbert & Heiner, 2006), as shown in detail in the supplementary material of this pa-211 per. 212





place affecting a controller. Meaning that T is a function of U_v



controller affecting a transition (flux). Meaning that parameters in E_T are a function of T

a splitter separating flux J into two contributions

Figure 5. Allowed connectivity between places, transitions and controllers, and a splitter in action. No other type of connection is possible.

To obtain the required budget equations, each place depicted in Figures 4 and 5 must correspond to the time variation of the quantity indicated in it. For instance, the

green place marked U_v represents the following part of a conservation equation:

$$\frac{dU_v}{dt} \tag{1}$$

with the quantity U_v being, for instance, the internal energy of a compartment of the HDSys. The differential operator can be changed for other operators, depending on the

type of equation we are writing, and, therefore a table defining which differential oper-

ator we are using is needed. From these rules we can represent a simple linear reservoir, as shown in Figure 6 on the left.



Figure 6. A simple linear reservoir (on the left) and a more complex example (on the right).

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In Figure 6 the flux J enters the place S_G , while the flux Q_G exits the same place. Therefore the budget is read as:

$$\frac{dS_g}{dt} = J(t) - Q_g(t) \tag{2}$$

Introducing another outgoing flux into the system, as shown on the right in Figure 6, the equation is modified to:

$$\frac{dS_g}{dt} = J(t) - Q_g(t) - E_T(t) \tag{3}$$

The action of the controller T on E_T remains hidden until we specify the mathematical form of the fluxes (transitions). This will be shown with the reference cases in the next section and mathematically formalized in section 8.

4 Casting the BST, HBV and Ard-Burn models into the EPN representation

Applying the rules introduced in section 3, we can now represent the three models of section 2 using EPN. We will present the details for the BST model, while we will be more concise for the others.

4.1 The BST model

As a result of the rules introduced in Section 3, the BST model, shown in Figure 1, can be represented using EPN as shown in Figure 7. It shows three coupled ODEs,

represented by three places, colored light blue, orange and dark red (colors chosen to be



Figure 7. Representation of the BST model (Birkel et al., 2011) using EPN. Compared with the original representation of Figure 1, this Figure contains less information, however, it is sufficient to write down the mass conservation equations for the system. The invisible reservoir is unnamed, since it is just the sum of Q_{low} and Q_{sat} and does not store water. As the legend shows, each color refers to a different conceptual-physical compartment through which the water flows. The outcomes from the splitter are named according to Table 1.

colorblind friendly, as better explained in the supplementary material). The small black bullets indicate quantities that should be measured and, therefore, assigned externally. A fourth, unnamed place has been added to highlight that measured data refers to the total flux, $Q_T = Q_{sat} + Q_{low}$, and not the two fluxes separately. This place is, in a sense, invisible because it does not introduce any ODE and its storage variation is always null; it has been left nameless and shown with dashed borders to reinforce this concept.

From the graph in Figure 7, the ruling equations are easily written as:

$$\frac{dS_{sup}(t)}{dt} = \underbrace{\alpha J(t)}_{J_l} - Q_1(t) - Q_R(t) \tag{4}$$

²⁴⁵ for the "sup" storage;

$$\frac{dS_{sat}(t)}{dt} = \underbrace{(1-\alpha)J(t)}_{J_T} + Q_1(t) - Q_{sat}(t) - E_T(t)$$
(5)

²⁴⁶ for the "sat" storage; and

$$\frac{dS_{low}(t)}{dt} = Q_R(t) - Q_{low}(t) \tag{6}$$

247 for the "low" storage.

Finally

$$0 = Q_T(t) - Q_{low}(t) - Q_{sat}(t)$$
(7)

In the BST model, there is one given (measured) input, precipitation J, which splits into J_l and J_r , and one given output, Q_T , each of which is marked with a small circle in Figure 7. One of the equations (the "orange" one, Eq. 5) contains a non-linear term, while the others are linear. Figure 7 is not sufficient to implement the model because its role (responsibility) is to identify the number of equations and to allow the reader to

- write the water budgets with unspecified fluxes. For complete information, two other elements are needed:
 - a **dictionary** giving the names of the symbols in the graphic (conveying their meaning), given in Table 1; and
 - an **expression table** giving mathematical completeness to the fluxes, presented in Table 2 . When there is a splitter, the corresponding flux is duplicated as necessary.

Expressions for places are not reported here since, by default, they associate any variable S_* to its time derivative dS_*/dt . However, in the most complex cases it is required to report them. Because the specification of fluxes usually introduces new variables, an extension to the dictionary may be necessary after writing the expression table. The substitution of the expressions in Table 2 into equations 4 to 7 gives the set of equations necessary to fully reproduce the model.

266	Table 1. Full dictionary associated to the EPN representation of the BST model (Birkel et
267	al., 2011). P stands for "parameter", F for "flux", SV for "state variable", V for "variable". [T]
268	stands for time units, $[L]$ for length units, $[E]$ for energy units. It contains the symbols present in
269	Figure 7 and also those implied by Table 2

	Symbol	Name	Type	Unit
	a	linear reservoir coefficient	Р	$[T^{-1}]$
	b	non-linear reservoir coefficient	Р	$[T^{-1}]$
	с	non-linear reservoir exponent	Р	[—]
	d	linear reservoir coefficient	Р	$[T^{-1}]$
	е	linear reservoir coefficient	Р	$[T^{-1}]$
	f	dimensional ET coefficient	Р	$[E^{-1}L^5]$
	$E_T(t)$	evapotranspiration	\mathbf{F}	$[L^3T^{-1}]$
	$J^{\bullet}(t)$	precipitation rate	\mathbf{F}	$[L^3T^{-1}]$
_	$J_l(t)$	precipitation rate going into S_{sup}	\mathbf{F}	$[L^3 T^{-1}]$
	$J_r(t)$	precipitation rate going into S_{sat}	\mathbf{F}	$[L^3T^{-1}]$
	$Q_1(t)$	discharge from the upper reservoir	\mathbf{F}	$[L^3T^{-1}]$
	$Q_{low}(t)$	discharge from the lower reservoir	\mathbf{F}	$[L^3 T^{-1}]$
	$Q_{sat}(t)$	discharge from the saturated reservoir	\mathbf{F}	$[L^3 T^{-1}]$
	$Q_R(t)$	recharge term of the lower reservoir	\mathbf{F}	$[L^3 T^{-1}]$
	$Q_T^o(T)$	total discharge at the outlet	\mathbf{F}	$[L^3 T^{-1}]$
	$R_n(t)$	net radiation	\mathbf{F}	$[EL^{-2}T^{-1}]$
	$S_{low}(t)$	storage in the lower reservoir	SV	$[L^3]$
	$S_{max}(t)$	maximum storage in the saturated reservoir	SV	$[L^3]$
	$S_{sat}(t)$	storage in the saturated reservoir	SV	$[L^3]$
	$S_{sup}(t)$	storage in the upper reservoir	SV	$[L^3]$
	t	time	V	
	α	partitioning coefficient	Р	[-]
	1.	· .		

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Table 2 clarifies the parameters of the model:

- J(t) is an external measured quantity (thus it is marked with a bullet, •);
- Only five parameters (a, b, c, d, e) are necessary since E_T is also assumed measured
- (as per original paper).

Table 2. Expression table associated to the EPN representation of the BST model(Birkel et al., 2011). Quantities marked with bullets represent measured quantities.

Flux	Name	Expression
$\overline{ET(t)}$	evapotranspiration	ET(t)
$J^{\bullet}(t)$	precipitation rate	•
$J_l(t)$	precipitation rate going into S_{sup}	$\alpha J^{\bullet}(t)$
$J_r(t)$	precipitation rate going into S_{sat}	$(1-\alpha)J^{\bullet}(t)$
$Q_{up}(t)$	discharge from the upper reservoir	$aS_{sup}(t)$
$Q_{low}(t)$	discharge from the lower reservoir	$dS_{low}(t)$
$Q_{sat}(t)$	discharge from the saturated reservoir	$bS_{sat}(t)^c$
$Q_R(t)$	recharge term of the lower reservoir	$eS_{up}(t)$
$Q_T^o(t)$	total discharge at the outlet	$Q_{sat} + Q_{low}$

4.2 The HBV model

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As another example, let us consider the EPN representation of the HBV model, shown in Figure 8. The HBV model was first shown in Figure 2 in Section 2. Tables A.1 and A.2 in Appendix A contain the associated dictionary and expression table.

The HBV model identifies four major compartments, snow (red), soil (yellow), ground-280 water (cvan) and surface waters (bright blue), as well as precipitation. It contains six 281 ODEs and, in contrast with the BST, it also contains a loop between SWE (snow wa-282 ter equivalent) and W_s (liquid water in snow). This loop implies that the output of liq-283 uid water from snow can refreeze and increase the amount the snow water equivalent from 284 which melted water derives and, by definition, adds a feedback to the system. This causes 285 some complications for the resolution of the model at the numerical level. In fact, parts 286 of graphs within loops have to be solved simultaneously with an iterative method (Car-287 rera, Holzbecher, Bonell, & Vasiliev, 2005; Patten, Higashi, & Burns, 1990), which usu-288 ally requires an overhead in computation proportional to the number of elements in the 280 loop. 290

A new feature appearing in the HBV model representation is the introduction of a controller. The triangle marked with T shows explicitly that temperature controls various fluxes, as made clear in Expression Table A.2: Actual Evapotranspiration, E_{act} , precipitation, P, melting rate of snow, M, and refreezing rate of the liquid water in the snowpack, R, are all controlled by temperature. For illustrative purposes, a fictitious dependence of T on S_{soil} has been added, with the scope of introducing controller dependent loops, which will be detailed in section 8.

4.3 The Loch Ard-Burn model

Finally, Figure 9 represents the Loch Ard-Burn model in Hrachowitz et al. (2013), 304 i.e. the model first shown in Figure 3 of Sector 2. The model has four major compart-305 ments: interception by vegetation (in green), an unsaturated reservoir (dark orange), a 306 fast reservoir (light blue), a slow reservoir (dark blue). In this case we, the Authors, have 307 preferred to identify the compartments with process names rather than locations and, 308 in a sense, this is also the choice in the perceptual model of the catchment. Compared 309 to the original representation, we have added three new reservoirs: the invisible S_O , and 310 X_F and X_S . S_O makes sense of the fluxes R_p and R_s that otherwise would both go from 311 S_{SU} into S_S without identifying them properly. In fact, the use of different kinds of blue 312 in the original representation in Figure 3 implies the existence of this reservoir. Hrachowitz 313 et al. (2013) introduced it to adjust the simulated water age to that measured with trac-314





ers. S_O does not accumulate water, implying that $R_S = R_O$, with a null net water budget exchange between S_O and S_{SU} reservoirs, but it mixes the younger water of the upper reservoirs with older waters to get the right water age at the budget. This trick was used before in Fenicia et al. (2010) and we shall not discuss it fully here.

In Hrachowitz et al. (2013) the discharges R_F and R_P from the unsaturated reservoirs seem to go to reservoirs S_F and S_S . However, these actually receive inputs R_F^* and R_P^* , which are the result of a convolution of R_F and R_P with some unit hydrographs. All of this implies the existence of additional reservoirs (places) to accommodate a water budget. For example, the discharges R_F and R_F^* are associated to the budget:

$$\frac{dX_F}{dt} = R_F - R_F^* \tag{8}$$

where the expression of the discharges is given in Table B.1 in Appendix B. In particular:

$$R_F^* = \int_0^t h_F(t - t_{in}) R_F(t_{in}) dt_{in}$$
(9)

where h_f is a instantaneous unit hydrograph whose expression is:

$$h_F(t) = \begin{cases} 1/2t/T_F^2 & 0 < t < T_F \\ 0 & otherwise \end{cases}$$
(10)

where t is time and T_F is a suitable parameter.

One might question whether this is the simplest modelling structure accounting for 321 tracer measurements and whether the place S_{SO} is necessary to have proper water ages 322 at the outlet. We merely observe that the representation in Figure 9 explicits this more 323 clearly than Figure 3. Besides, Figure 3 ignores the existence of the discharge R_U , which 324 is necessary to preserve the mass budget of the unsaturated reservoir S_{SU} . It is worth 325 noting how the inclusion of controllers in the EPN representation shows clearly the im-326 portance of potential evapotranspiration E_p and the C_E parameter (a function of stor-327 ages S_{SU} and S_F) on evapotranspiration; otherwise, this would only be apparent by a 328 careful inspection of the flux expressions. The Dictionary and Tables for the Ard-Burn 329 model are presented in Appendix B. 330

³³⁵ 5 Use of Petri Nets for interpreting field work

EPN can be used during the "perceptual phase" of research that moves from ex-341 perimental evidence to the construction of an appropriate numerical model of a catch-342 ment. This can be done either according to the strategies defined in Fenicia and Kavet-343 ski (2011) and Clark et al. (2015), or with a more qualitative procedure, like the one we follow here, which represents just one practical application of EPN's functionalities. As 345 an example, we can take the description of the Maimai catchment (Gabrielli et al., 2018), 346 which is probably among the most widely studied small catchments in the world. The 347 dynamics of the catchment is described as: "Catchment storage is formed by two sharply contrasting and distinct hydrological units: shallow, young soil storage, and deep, much 349 older bedrock groundwater". Therefore, there are at least two storage reservoirs. The 350 description then continues: "This storage pairing produces a bimodal, seasonal stream-351 water". This means that streams are a third reservoir that collect water from the other 352 two, the soil and groundwater reservoirs. It then states that during the summer months 353 there is evapotranspiration, E_T , and that it is an important term of the water budget. 354 In a conceptual model E_T can only come from the soil reservoir. The groundwater reser-355 voir contributes to surface waters and downstream storage. A proper description of the 356 catchment should also include the effects of interception and evaporation from the canopy; 357 however, for simplicity, these are not taken into account here. 358

From this description, then, it seems that the perceptual model can be instantiated with two EPN places, which correspond to a set of two main ordinary differential



Figure 9. EPN representation of the Ard-Burn model, corrected for proper water age tracking. It has seven main water budget equations, derived from an accurate reading of Hrachowitz et al. (2013). The red dotted reservoir, S_O , is added to account properly for tracer history. The bluish reservoirs account for lag times from $S_{SU} \rightarrow S_F$ and $S_{SU} \rightarrow S_S$.

equations, as shown in Figure 10. Because of its similarity with the system proposed by Kirchner (2016), we have used the names introduced in that paper, with the exception of evapotranspiration, E_{T_s} , and percolation, R_l , which we have added.

Another reservoir can be added to account for surface water storage where groundwater and soil water mix. This reservoir is where the fluxes \check{L} and Q_l are summed and, as such, it is an invisible place. The dictionary for this system is presented in Table 3. To understand how to write the tentative equations for such a system, we need to further clarify the semantics of the graph, i.e. we need to make the mathematical structure of the fluxes explicit. For this one can find inspiration in Kirchner (2016) but we do not pursue it further here.

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6 Modeling Hydrology as an Earth System Science

The HDSys are open dynamical systems that exchange water and energy with their surroundings. They are non-linear and usually non-autonomous, they have non-trivial time-dependent properties and, being open systems, their future inputs are unknown. Therefore, they differ from the dynamical systems treated in other disciplines where, for instance, forcings can be written as periodic functions (a typical example in textbooks is Strogatz (1994)).

One of the contemporary directions of hydrological research is to investigate HDSys as part of the larger Earth system science, which includes, among others, the energy and carbon cycles. Thus, the hydrological cycle becomes part of a broader living environment



Figure 10. EPN representation of the Maimai catchment according to our reconstruction. It has two main reservoirs, a soil reservoir, S_u , and a groundwater reservoir, S_l . There is also a surface water reservoir, S_{sup} , where soil waters and groundwater mix without any delay (so it is an invisible place). In Gabrielli et al. (2018) soil water fluxes and groundwater fluxes were measured separately and, therefore, we mark them with a black bullet.

Table 3. Dictionary for the Maimai catchment model. F indicates "flux"; SV "state variable";
 V "variable". Quantities marked with bullets represent measured quantities.

Symbol	Name	Type	Units
$\overline{E_{T_s}(t)}$	Evapotranspiration from the soil reservoir	F	$[L T^{-1}]$
$\check{L}(t)$	Discharge from soil	F	$[L T^{-1}]$
$\hat{L}(t)$	Recharge to groundwater	F	$[L T^{-1}]$
$P^{\bullet}(t)$	Precipitation	F	$[L T^{-1}]$
$Q_l(t)$	Discharge from groundwater	F	$[L T^{-1}]$
$Q_S^o(t)$	Total discharge	\mathbf{F}	$[L T^{-1}]$
$R_l(t)$	Percolation to a deeper aquifer	F	$[L T^{-1}]$
$S_l(t)$	Storage in the groundwater reservoir	SV	[L]
$S_u(t)$	Storage in the soil reservoir	SV	[L]
t	time	V	[T]

- tems are not passive spectators of hydrological events but co-evolve with hydrology (H. G. Savenije
- ³⁸⁵ & Hrachowitz, 2017). According to this concept, ecosystems control the hydrological cy-

that feeds back on itself (H. H. G. Savenije & Hrachowitz, 2017; Zehe et al., 2014). Ecosys-

cle (and vice versa, of course). To be able to represent such complexities, we have to ensure that EPN can represent the energy budget and vegetation growth just as well as
it represents the water budget. For these aspects, clearly, the usual representation of a
model as a complex of reservoirs falls short.



Figure 11. Coupled energy and water budgets. The graphic notation is enriched with the 391 addition of a new type of arc (dotted segments ending in empty squares). These arcs connect the 392 same variables present in both budgets. In this case, the J's are input, while the E_T 's and Q_G 's 393 are unknown variables that must be solved simultaneously in both budgets. Because E_T depends 394 on radiation, a controller exiting from the U_G place is added to reveal this further influence of the 395 energy budget on the water budget. Other controllers of the system can be the leaf area index, 396 LAI, which controls radiation and evapotranspiration, and hydraulic conductivity, K_S , which can 397 be thought to influence flow Q_G . 398

Rarely has the energy budget been present in hydrological models so far. Current 300 studies, covering the whole set of hydrological fluxes (e.g. Abera, Formetta, Brocca, and 400 Rigon (2017); Kuppel, Tetzlaff, Maneta, and Soulsby (2018)), require that both the wa-401 ter and energy budgets be solved. To describe this coupling we use the simple example 402 shown in Figure 11(left), referring to a hillslope water budget, with the the associated 403 energy budget also shown in Figure 11(right). To distinguish between the budgets, we 404 used a further graphical stratagem in the figure and colored the background light pas-405 tel blue for the water budget and light pastel red for the energy budget). 406

The dictionary associated to the graph in Figure 11 (left) is in Table 4 and the budget can be deduced to be:

$$\frac{dS_g(t)}{dt} = J(t) - E_T - Q_g \tag{11}$$

⁴⁰⁷ The Expression Table is not needed at present and has been omitted.

408	Table 4.	Dictionary d relative	to Figure 11.	The underscoring (\cdot)) represents the internal en-
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ergy acquired or lost through mass exchanges. 400

	Concernant States			
	Symbol	Name	Type	Unit
	[E]	energy per unit area	-	[E]
	$E_T(t)$	evapotranspiration	F	$[LT^{-1}]$
	$\underline{E}_{T}(t)$	evapotranspiration energy content	\mathbf{F}	$[EL^{-2}]$
	H	sensible heat	\mathbf{F}	$[ET^{-1}]$
	$J^{\bullet}(t)$	precipitation rate	\mathbf{F}	$[LT^{-1}]$
ь	$J^{\bullet}(t)$	precipitation energy content	\mathbf{F}	$[ET^{-1}]$
~	$\overline{Q_q(t)}$	discharge	\mathbf{F}	$[L^3T^{-1}]$
	$Q_{a}(t)$	discharge internal energy	F	$[ET^{-1}]$
	$\overline{R}^{g}\downarrow$	incoming radiation	F	$[ET^{-1}]$
	$R\uparrow$	outgoing radiation	\mathbf{F}	$[ET^{-1}]$
	$S_q(t)$	water storage	SV	$[L^3]$
	t	time	V	[T]
	U_g	internal energy	SV	[E]

In Figure 11(right), one can observe that the internal energy of the control volume contains one energy flux for each water flux present in the water budget. In fact, each mass flux has an associated internal energy, conveniently represented as enthalpy per unit mass, which flows in or out when mass is acquired or lost by the control volume. Thus, for instance, given the rainfall J, the corresponding enthalpy flux is $\underline{J} = \rho_w h_w J$, where ρ_w is the water density in the volume, and h_w is the water enthalpy per unit mass. In short, many variables are common to both budgets, i.e. they are shared by the budgets and must satisfy both of them. These variables are joined by a new type of arc, a dotted segment capped with empty squares. In addition to these variables, in the energy budget we have to account for the radiation budget, written here as the budget of incoming $R \downarrow$ and outgoing, $R \uparrow$ radiation associated to the place U_G . Latent heat is accounted for as evapotranspiration multiplied by the latent heat (enthalpy) of vaporization. Finally, the energy flux due to thermal energy exchange by convection (sensible heat), flux H, is taken into account. The resulting energy budget equation is:

$$\frac{dU_G}{dt} = \underline{J} + R \downarrow -R \uparrow -\underline{E}_T - H - \underline{Q}_G \tag{12}$$

410 411 Table

Table 5.	Expression	table E	relative	to the	energy	exchange	model	presented	in	Figure	11	on
the right.												

Symbol	Name	Unit
$E_T(t)$	evapotranspiration	$[ET^{-1}L^{-2}]$
$\overline{H(t)}$	thermal convective flux	$[EL^{-2}T^{-1}]$
J(t)	precipitation rate	$[ET^{-1}L^{-2}]$
$\overline{\mathcal{J}_g}$	thermal conduction losses to the ground	$[ET^{-1}L^{-2}]$
$Q_g(t)$	discharge	$[ET^{-1}L^{-2}]$
$\overline{R_n(t)}$	Net Radiation	$[EL^{-2}T^{-1}]$
$U_g(t)$	internal energy storage per unit area	$[EL^{-2}]$

Furthermore, hydraulic conductivity, K_S , is thought to control the water flux, Q_G , while the Leaf Area Index (LAI) controls evapotranspiration and radiation response of the system (through long wave radiation fluxes). Admittedly, some simplification have been made when coupling the water budget with the energy budget, however, the procedure is quite general and can be used for more complicated cases.



Figure 12. The simple vegetation growth model presented in Montaldo et al. (2005). It consists of three coupled ODEs, which account for aboveground, below ground and dead vegetation.

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The interactions of water budget with the ecosystem can also be represented with EPN. As an example, we use a simple vegetation growth model presented in Montaldo et al. (2005) and further developed in Della Chiesa et al. (2014). The model consists of three ODEs, for above ground vegetation, B_g , roots, B_r , and dead material, B_d :

$$\frac{dB_g}{dt} = a_a P_g + T_{ra} - R_g - S_g \tag{13}$$

where B_g is the mass of the green above ground biomass, P_g is the gross photosynthesis, a_a is the allocation partition coefficient to shoots, T_{ra} is the translocation of carbohydrates from the roots to the living above ground biomass, R_g is the respiration of the above ground biomass, and S_g the senescence of the above ground biomass.

$$\frac{dB_r}{dt} = a_r P_g - T_{ra} - R_r - S_r \tag{14}$$

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where B_r is the living root biomass, a_r ($a_r + a_a = 1$) is the allocation partition coefficient to roots, R_r is the respiration from roots, S_r the senescence of roots.

$$\frac{dB_d}{dt} = S_g - L_a \tag{15}$$

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where B_d are the standing dead, S_g is the senescence of aboveground biomass and L_A is the litter fall. All of these quantities are described in the dictionary in Table 6 and are represented by the EPN in Figure 12. This model is presented to show how vegeta-

D	Table 6.	Dictionary relative t	o the model o	f vegetation	growth in	Montaldo e	et al.	(2005)	and
1	illustrated	in Figure 12.							

Symbol	Name	Type	Unit
$\overline{a_a}$	allocation partition coefficient for aboveground biomass	Р	[-]
a_r	allocation partition coefficient for root compartments	Р	[-]
B_d	standing dead biomass	SV	$[M L^{-2}]$
B_q	green aboveground biomass	SV	$\left[M L^{-2} \right]$
B_r	living root biomass	SV	$[M L^{-2}]$
L_a	litter fall	\mathbf{F}	$[M L^{-2} T^{-1}]$
P_{q}	gross photosynthesis	\mathbf{F}	$[M L^{-2} T^{-1}]$
$\vec{R_q}$	transpiration from aboveground biomass	\mathbf{F}	$[M L^{-2} T^{-1}]$
R_r	transpiration from root biomass	\mathbf{F}	$[M L^{-2} T^{-1}]$
S_{q}	senescence of the aboveground biomass	\mathbf{F}	$[M L^{-2} T^{-1}]$
S_r	senescence of the root biomass	F	$M L^{-2} T^{-1}$
T_{ra}	translocation of carbohydrates from roots to the above ground biomass	F	$[M L^{-2} T^{-1}]$

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tion can interact with the hydrological cycle, an aspect that can be fully revealed only

through an expression table. For the sake of simplicity, Table 7 does not contain the com-

⁴²⁸ plete mathematical expressions, which are fully discussed in Della Chiesa et al. (2014);

- ⁴²⁹ Montaldo et al. (2005), but it does provide the variable dependence needed to produce the h-connections between the vegetation model and the water and energy budgets. The
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 Table 7. Expression Table relative to the model of vegetation growth in Figure 13.

Symbol	Name	Expression
P_g	gross photosynthesis	$P_g(\Delta CO_2, r_a, r_c)$

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interesting fact is that, through parameters like the LAI, the aboveground vegetation controls evapotranspiration and radiation, while roots are thought to control the hydraulic conductivity, K_S . Photosynthesis feeds a vegetation system and is controlled by variables such as temperature, T, photosynthetic active radiation (here made dependent on the energy budget), and soil water content, θ . All of this is represented in Figure 13 and is discussed in the next section.

437 7 Discussion

⁴³⁸ While the graphs of the water budget, energy budget and vegetation growth are themselves direct, acyclic graphs, the whole coupled graph, inclusive of h-wiring, shows loops, like the one between $U_G \to T \to P_g \to B_r \to LAI \to R \downarrow \to U_G$, that depict a feedback. Therefore, to really understand the interactions between the three dynamical systems graphically, we have to use h-wiring as we do in Figure 13. Notably, while the water budget can be represented with traditional reservoirs, the traditional graphics fall short in representing the other budgets.

Figure 13 is only for demonstration purposes and, as such, the connections shown 447 are hypothetical. As we have not implemented and tested such a model, the relations 448 presented are based on educated guesses. The quantities that appear in the h-wiring net-449 work are constraints on the dynamical model parameters and work as valves that reg-450 ulate the fluxes. Without h-wiring, the connections between sub-models are not evident. 451 Although, the flux connections (f-wiring) alone are sufficient to write the correct ODEs 452 in their completeness, including feedback loops, once complemented by the appropriate 453 expression tables. When H. G. Savenije and Hrachowitz (2017) write, "The most impor-454 tant active agent in catchments is the ecosystem. [...]. Ecosystems do this in the most 455 efficient way, establishing a continuous, ever-evolving feedback loop with the landscape 456 and climatic drivers", they refer to the ability of ecosystems, represented in Figure 13 457 by the bottom set of ODEs, to control the water cycle. The Figure shows how this hap-458 pens through the action of controllers that link vegetation to both the water and energy 459 cycles. We do not know yet if the system devised includes the right properties to obtain 460 the dynamical richness desired. To get an answer one should look towards system and 461 control theories. These (Kalman, 1959) offer more than fifty years' worth of literature to help deal properly with interacting systems. In fact, one pivotal concept in system 463 and control theories is **controllability**, i.e. the possibility that a system that has drifted 464 into an undesirable state can be steered back to another desirable one. Linear theory (Willems, 465 2007) contains theorems and tools (Kalilath, 1980; Luenberger, 1979; Sontag, 1998) that can assess controllability precisely but, unfortunately, our HDSys is not linear and, at 467 first sight, our controllers do not seem to fit the concept of **actuators**, the agents that 468 perform the control. 469

To treat non-linearities more completely, more sophisticated analyses are needed, 470 (Liu & Barabasi, 2016). Fortunately, a lot has been accomplished since the 1970s (Cor-471 nelius & Kath, 2013; Haynes & Hermes, 1970; Hermann & Krener, 1977). Great strides 472 have been made, both from an analytical point of view and from a graph theory point 473 of view, (Liu & Barabasi, 2016; Yamada & Foulds, 1990). Notably, the latter results are 474 directly interpretable by using the EPN presented here, though an exploitation of these 475 possibilities goes beyond the scopes of this paper. However, it should be noted that any 476 graphical representation that does not contain fluxes in explicit form (i.e. as nodes of 477 the graphs) and h-wiring, brings to a scanty graphical representations of the dynamics 478 and, as a consequence, to incorrect graphical analyses. 479

The discussion so far has been referred to a single spatial unit or HRU. If a catchment is divided into various parts, the EPN of the single spatial units can be joined to obtain the integral distributed view of the basin. For illustrative purposes, in this paper we use a simple catchment partition based on the identification of subcatchments, as shown in Figure 14.

In the example case, the basin is subdivided in 5 HRUs (Figure 14, top left), which 485 have been derived by dividing the river network into five links C_1 to C_5 . It is assumed 486 that the external fluxes to the HRU are rainfall J_i in input, and evapotranspiration E_{Ti} 487 and discharges Q_i $(i \in \{1, ..., 5\})$ in output. Each HRU flows into a channel stream, for 488 instance, the HRU of area A_4 flows into C_4 and subsequently to C_5 . The complete net-489 work of interactions can be represented as in Figure 14. A black frame marking some 490 of the external places indicates that they are actually *compound places*. These can be 491 expanded by using embedded models, like those shown in the Figures of the previous sec-492 tions or some generalization of the more complex model of Figure 13. The HBV model 493 (Seibert & Vis, 2012) is meant to be just such a model: the HBV structure presented in Figures 2 and 8 can be used for any sub-catchment of the basin analyzed. 495



Figure 13. Representation of the water and energy budgets coupled with the vegetation
dynamic model. The coupling happens entirely through h-wiring



Figure 14. A small river networks with 5 HRUs (top left) and the corresponding EPN. There
is a black frame marking some of the places to indicate that they are compound places; these
should be expanded to reveal the full structure of the system.

Figure 14 exploits the compositionality of EPN and shows how it can be used to represent any river network. Semi-distributed modeling can become very complex and even have heterogeneous elements in each compound node. It is not a matter for this paper to discuss when it becomes too complicated to be reasonably useful. The scopes of our representation is to make the model structures presented as clear as possible and, eventually, to exemplify it. To pursue the latter task, it is helpful to translate the graphs into mathematics, as we do in the next section.

8 A formal mathematical treatment of EPN

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So far we have treated the graphics and their relation to mathematics in a conver-507 sational way. However, these relations can be more precisely stated by providing a set 508 of definitions for the entities appearing in EPN, which the reader can find below. The 509 definitions have the advantage of formalizing the topology of the models by introduc-510 ing appropriate adjacency and incidence matrixes. These matrixes, in turn, reveal that 511 the structure of the hydrological dynamical system can be studied objectively using tech-512 niques derived by algebraic topology (Fiedler, 1973) and, as mentioned in the previous 513 sections, already used in other fields Our definitions (any item marked with a bullet, \bullet) 514 expand the notation introduced in Navarro-Gutiérrez, Ramírez-Treviño, and Gómez-Gutiérrez 515 (2013) and are modified as suggested by Baez and Pollard (2017). To exemplify them, 516 we will refer to that part of the HBV model that has been framed in black in Figure 8. 517

8.1 The topology of a HDSys

• $\mathcal{P} = \{p_1, \cdot, p_n\}$ is the set of *n* places (reservoirs). In our graphical notation, they are identified by *n* circles. In the HBV example, $\mathcal{P} = \{SWE, W_S, S_{soil}\}$.

• $\mathcal{T} = \{t_1, \cdot, t_l\}$ is the set of l transitions (fluxes). Graphically, they are represented by l squares. In the HBV example, $\mathcal{T} = \{M, R, M_d, F, E_{act}, P\}$

In EPN the relationships between these two types of nodes (i.e. places and transitions) can be expressed with two incidence matrices.

> A^- is the incidence matrix that represents the connections from places to transitions, i.e. it is an $n \times l$ matrix, where the element (i, j) is marked with 1 if place *i* outputs to transition *j* and otherwise it is 0. In our graphical notation the connections are shown with oriented arcs joining the appropriate couple (p_i, t_j) . With respect to the HBV example, A^- is shown in table 8.

_	$\mid R$	M	M_d	F	E_{act}	P
	1					
SWE	0	1	0	0	0	0
W_s	1	0	1	0	0	0
Ssoil	0	0	0	1	1	0
Table	8	4 - m	atriv fo	r the	HRV	won

Table 8. A^- matrix for the HBV example. P is an input and has no places connecting to it, therefore, its column does not contain any 1s.

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• A^+ is the incidence matrix that represents connections from transitions to places, i.e. it is an $l \times n$ matrix, where the element (k, m) is marked with 1 if transition k is an input to place m; otherwise it is 0. Graphically the connections are oriented arcs joining (t_k, p_m) for the appropriate k and m. The A^+ matrix relative to the HBV example is shown in Table 9

_	$\mid SWE$	W_S	S_{Soil}
\overline{R}	1	0	0
M	0	1	0
M_d	0	0	1
F	0	0	0
E_{act}	0	0	0
P	1	0	1

Table 9. The A^+ matrix relative to the HBV model. F and E_{act} are outputs of the whole system and therefore their rows contain only 0s.

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- There are two possible products of the incidence matrices, A^+ and A^- , both of which result in a square matrix:
- $A = A^{-} \cdot A^{+}$ is the $(n \times n)$ adjacency matrix that identifies the connections between places. The A matrix for the HBV model is shown below in Table 10 • $\tilde{A} = A^{+} \cdot A^{-}$ is the $(l \times l)$ adjacency matrix that identifies the connection between transitions. The \tilde{A} for the HBV example is presented in Table 11

Transitions and places, and their relationships as expressed in incidence and adjacency matrices, can be used to represent the ODE system of any budget (mass, energy, momentum).

		$\mid S$	WE	W_S	S_S	oil	
	SWE	0		1	0		
	W_S	1		0	1		
	S_{Soil}	0		0	0		
541	Table	e 10.	The	e A ma	trix f	for the l	HBV example. The anti-diagonal 1s reveal the presence of a
542	loop.	1					
	P	2					
	-	1					
	-	$\mid R$	M	M_d	F	E_{act}	Р
	\overline{R}	0	1	0	0	0	0
	\overline{M}	1	0	1	0	0	0
-	M_d	0	0	0	1	1	0
	F	0	0	0	0	0	0
	E_{act}	0	0	0	0	0	0
	P	0	1	0	1	1	0

Table 11. The \tilde{A} adjacency matrix for transitions with respect to the HBV example. It reveals the connections between fluxes

Starting from any one of the places (circles), transitions (squares) in the graphic and:

> • following the arcs we get a **causal path**. When two variables are connected by an arc, the upstream entity is said to cause the downstream one. Therefore a transition is caused by the upstream place and a place by the upstream fluxes. Causality is inherited, in that all upstream variables have causal influence on downstream ones.

However, the resulting EPN, do not show the feedbacks between state variables 559 completely because some of these can be hidden in the flux expressions. Therefore, to 560 provide a more complete visual representation of the causal relationships between vari-561 ables, we have introduced the concept of controllers. Controllers are a function of a state 562 variable (originated in a place) that contribute in the flux expressions of one or more tran-563 sitions. They are explicitly represented by triangles in the graph. A rectangular incidence 564 matrix B, of dimensions $(n \times l)$, indicates the places that are connected to transitions 565 via controllers. The resulting web of interactions is called hidden wiring or **h-wiring**. 566 B can be split into two matrices (as was the case for A). 567

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If:

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- $C = \{c_1, \cdot, c_m\}$ is the set of controllers. In the HBV example there is just one controller, T, the temperature, therefore $C = \{T\}$.
- B^- is the incidence matrix representing the connections from places to controllers. It is an $n \times m$ matrix with the non-null element (i,j) set to 1 if place *i* is connected to controller *j*. Graphically, oriented dashed arcs are used to connect circles to triangles. The B^- matrix of the HBV example is represented in Table 12.
- B^+ is the incidence matrix $(m \times l)$ between controllers and transitions. Graphically the connection between controllers and transitions are represented by oriented dashed arcs between triangles and squares. The usual example from the HBV model reads as in Table 13.

581 Then:



	$SWE W_s S_{Soil}$
	$\overline{SWE \mid 0}$ 1 0
	W_s 1 0 1
	S_{Soil} 2 1 1
595	Table 16. The D matrix for the HBV example. It represents all the connections between
596	places, either mediated by fluxes or by h-wiring.
604	places to controllers, and from controllers to transitions, that we call the topol-
605	ogy of the EPN
005	
606	Models with the same topology can have different fluxes and state variables.
	1 00
607	8.2. The semantics of a HDSvs
007	5.2 The semances of a HDSys
608	The semantics provide all the information needed to complete the equations of
609	a given system on the basis of its topology. Let us define the semantics as follows.
	Lat
610	Let.
611	• D be the dictionary or lexicon of a model. It associates each symbol in the topol-
612	ogy to its meaning (and other information such as units and the role of the vari-
613	able). Various examples were given in the previous sections, such as Tables 1 above
614	and Table A.1 below.
615	• S be the set of expressions for places, associating to each place its mathematical
616	operator (in this paper the default expression for a place is the time derivative of
617	the state variable);
618	• E be the set of expressions for fluxes, associating to each flux its algebraic form.
619	Examples are given in Tables 2 and B;
620	• C be the set of expressions that define controllers as functions of state variables.
621	Table B3 is an example for the Ard-Burn example.
622	Then,
623	• the semantics of an EPN is the quadruple: $\mathcal{Y} = (D, E, C, S)$
<i></i>	Finally
624	r many,
625	• The pair $\mathcal{M} = (\mathcal{X}, \mathcal{V})$ (topology and semantics) fully defines a HDSvs.
020	
<i>cac</i>	8.3 A definition of hydrological dynamical systems
626	6.6 A demittion of hydrological dynamical systems
627	Some further definitions can be useful in understanding the nature of the model
628	$\mathcal{M}.$
c	The set:
629	
630	• $s_i = \{p_i A_{ii}^- > 0\}$ is said to be the preset (or the set of sources) of the tran-
631	sition t_i : In Table 8 this set can be deduced from the non-null terms in the columns.
632	• $o_i = \{p_i A_{i,i}^+ > 0\}$ is said to be the postset or the set of targets of transition
633	t_i . In Table 9 they are the non-null terms in any row.
	v v
634	Then,



Models are compositional in the sense that, given two model \mathcal{M} and \mathcal{M}' , we say that

 \mathcal{M} and \mathcal{M}' can be composed if at least one output of one model coincides with one input of the other.

However, models can also be composed by sharing controllers. For instance, a model of 661 the energy budget can provide the temperature T, which is a controller of the model HBV. 662 Thence, the energy budget not only constrains the behavior of HBV but can also be com-663 posed with it. We can say that, given two models \mathcal{M} and \mathcal{M}' , they can be composed: 664

• by sharing fluxes (**f-wiring**); or

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• by sharing controllers (**h-wiring**)

This means that our models, and their representations, have at least two associative prop-667

erties that can be used to obtain arbitrarily complicated models. The use of h-wiring is 668

(16)

(17)

- only fully possible with a Petri net type of representation because in other graphical sys-
- tems, fluxes (transitions) do not have the graphical status of nodes.

Articl

9 Conclusions

In this paper we introduced an extension of Petri Nets to describe lumped hydrological models and make evident that they are part of the great family of dynamical systems and/or compartmental models. The EPN representation:

- is adequate to describe any lumped hydrological system and the interactions be-675 tween the hydrological, energy and carbon cycles, which form the basis for the mod-676 elling of Earth system interactions. 677 • standardizes the way to represent hydrological models and interactions; 678 • streamlines the process of documenting hydrological models; 679 • facilitates user comprehension of eco-hydrological interactions (number of places 680 corresponds to the number of equations, number of transitions to the number of 681 fluxes, and number of controllers to the number of constraints imposed on the fluxes); 682 • can be used to organize process interactions hierarchically, even when the math-683 ematical flux expressions are not set; 684 allows for an easy comparison of model structures in terms of topology and seman-685 tics (via specific expression of fluxes and constraints); 686 visually represents feedback loops between subcomponents, even those implied by 687 non-linear terms, that are hidden in other treatments of the subject; 688 provides a complete visual representation of the causal relation between variables 689 used in models; 690 helps to understand lumped models as systems of systems of ODEs that can be 691 composed to form larger systems; 692 • builds a bridge with analysis techniques developed in mathematics or other dis-693 ciplines, such as theoretical biology, neuroscience and computer science; 694 • hints how results from linear and non-linear Systems and Control theory can be 695 used to gain insight into hydrological processes and evaluate the control exerted 696 by ecosystems on hydrology and by hydrology on ecosystems. 697
- At the same time, being general, EPN can be easily used in other disciplines, such as ecology, chemistry, biology and population dynamics.

A Dictionaries and Expression table for the HBV model

In this Appendix we report the Dictionary and the Expression table for the HBV model. The information presented, together with the EPN, allows one to write the dynamical system that corresponds to the HBV model with confidence.

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Table A.1. Dictionary for the HBV model Seibert and Vis (2012). P type stands for "parameter"; F for "flux"; SV for "state variable"; C for controller; V for independent variable

Symbol	Name	Type	Units
E_{act}	actual evapotranspiration	F	$[L T^{-1}]$
E_{pot}	potential evapotranspiration	\mathbf{F}	$[L T^{-1}]$
$E_{POT,M}$	long term mean potential evapotranspiration	Р	$[L T^{-1}]$
F(t)	flux of water to the upper reservoir	\mathbf{F}	$[L T^{-1}]$
M	rate of snow melting	\mathbf{F}	$[L T^{-1}]$
M_d	release of liquid water from snow	\mathbf{F}	$[L T^{-1}]$
P^{\bullet}	precipitation	\mathbf{F}	$[L T^{-1}]$
P_{BETA}	exponent in flux to upper zone	Р	[-]
P_{CFMAX}	degree-day factor in snow melting	Р	$[L T^{-1}]$
P_{CFR}	proportion of water refreezing	Р	[-]
P_{CET}	parameter in defining E_{POT}	Р	[T-1]
P_{FC}	maximum value of soil storage	Р	[L]
P_{K0}	parameter in estimation of flux out of upper zone	Р	$[T^{-1}]$
P_{K1}	parameter in estimation of flux out of upper zone	Р	$[T^{-1}]$
P_{K2}	parameter in estimation of flux out of LZ	Р	$[T^{-1}]$
P_{LT}	parameter: entering in evaporation estimation	Р	[-]
P_{perc}	percolation to groundwater	\mathbf{F}	$[L T^{-1}]$
\dot{P}_{MAXBAS}	parameter: in definition of $c(i)$	Р	[-]
P_{TT}	threshold parameter for melting activation	Р	[T]
Q_{GW1}	runoff from the upper zone to the surface waters	\mathbf{F}	$[L T^{-1}]$
Q_{GW2}	groundwater flow	\mathbf{F}	$L T^{-1}$
Q_{sim}	river network discharge	\mathbf{F}	$[L T^{-1}]$
R	rate of liquid water refreezing	\mathbf{F}	$[L T^{-1}]$
S_{soil}	water in soil/root zone	SV	[L]
S_{LZ}	groundwater storage	SV	[L]
S_R	runoff storage	SV	Ĺ
S_{UZ}	water Storage in the upper zone	SV	Ĺ
SWE	Snow Water Equivalent	SV	[L]
T^{\bullet}	temperature	С	[T]
T_M	long-term average temperature	Р	[T]
W_S	liquid water in snow	SV	Ĺ

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The expressions in Table A.2 are quite long, given our desire to respect the names used in the paper Seibert and Vis (2012); we are forced, therefore, to introduce the ancillary table A.3 that contains the missing sub-expressions. Once sub-expressions are sub-710 stituted into their corresponding variable, the complete form of the fluxes is obtained. 711

B Dictionary and Expression table for the Loch Ard-Burn model

Here we present the dictionary and the expression table for the Loch Ard-Burn model. 714 Notwithstanding the apparent simplicity of Figure 3, the model becomes quite compli-715 cated when complete information is provided. 716

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Table A.2. Expression table for HBV model. The flux expressions are quite long and, there-

⁷⁰⁷ fore, some ancillary quantities are defined in table A.3

Flux	Name	Expression
E_{act}	actual evapotranspiration	$E_{pot} \min\left(\frac{S_{soil}(t)}{P_{FC}P_{LT}}, 1\right)$
F(t)	flux of water to the upper reservoir	$I(t) \left(\frac{S_{soil}}{P_{EC}}\right)^{P_{BETA}}$
Μ	rate of snow melting	$P_{CFMAX}(T(t) - P_{TT})$
M_d	release of liquid water from snow	M-R
P	precipitation	•
P_{perc}	percolation to groundwater	
\dot{Q}_{GW1}	runoff from the upper zone	$P_{K2}S_{LZ}$
	to the surface waters	
Q_{GW2}	groundwater flow	$P_{K0} \max (S_{UZ} - P_{UZL}, 0) + P_{K1} S_{UZ}$
Q_{sim}	river network discharge	$\sum_{i=1}^{P_{MAXBAS}} c(i) (Q_{GW1}(t-i+1)+$
		$\overline{Q}_{GW2}(t-i+1))$
R	rate of liquid water refreezing	$P_{CFR}P_{CFMAX}(P_{TT}-T(t))$

 Table A.3.
 Table of ancillary variables in the HBV model Expression Table

Variable	Name	Expression
$\overline{ egin{array}{c} c(i) \ E_{pot} \ I(t) \end{array} }$	ancillary variable in Q_{sim} potential evapotranspiration sum of snow melt and precipitation	$ \begin{array}{c} \int_{i-1}^{i} \frac{2}{P_{MAXBAS}} - \left u - \frac{P_{MAXBAS}}{2} \right \frac{4}{P_{MAXBAS}^{2}} du \\ (1 + P_{CET}(T(t) - T_{M}))E_{pot,M}) \\ S_{soil} + M_{D} \end{array} $

717	Table B.1.	Dictionary for the Loch Ard-Burn model. Most of the nomenclature derives from
718	Hrachowitz et	al. (2013). However, in that paper, the S_O reservoir is not drawn and, in our opin-
719	ion, some sym	bols were not named properly; the underlined words represent more appropriate

names, in our opinion.

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Symbol	Name	Type	Units
C_E	partition coefficient between evapotranspirations	С	[-]
C_R	partition coefficients between runoff types	Р	[-]
E_p	potential Evapotranspiration	\mathbf{F}	$[L T^{-1}]$
$\hat{E_{SI}}$	evaporation from vegetation	F	$[L T^{-1}]$
E_{SU}	transpiration from unsaturated reservoir	F	$[L T^{-1}]$
E_{SF}	transpiration from fast responding reservoir	F	$[L T^{-1}]$
h_F	response time distribution for X_F reservoir	[]	$[T^{-1}]$
h_S	response time distribution for X_S reservoir	Ū	$[T^{-1}]$
K_F	storage coefficient of fast reservoir	Р	$[T^{-1}]$
K_S	storage coefficient of slow reservoir	Р	$[T^{-1}]$
I_{max}	maximum interception	Р	[L]
L_p	transpiration threshold	Р	[-]
P_{max}	percolation capacity	Р	$[L T^{-1}]$
P_R^{\bullet}	rainfall	F	$[L T^{-1}]$
P_{TF}	throughfall	F	$[L T^{-1}]$
Q_{SF}	runoff from fast reservoir	F	$[L T^{-1}]$
Q_{OF}	overland flow	F	$[L T^{-1}]$
Q_{SS}	runoff from slow reservoir	F	$[L T^{-1}]$
R_F	recharge of fast reservoir	F	$[L T^{-1}]$
R_O	flux from hidden old water reservoir to unsaturated zone	F	$[L T^{-1}]$

Symbol	Name	Type	Units
R_P	preferential recharge of slow reservoir	F	$[L T^{-1}]$
R_S	recharge of <u>old water</u> reservoir	\mathbf{F}	$[L] T^{-1}$
R_U	percolation from the unsaturated reservoir	\mathbf{F}	$[L T^{-1}]$
S_I	intercepted storage	SV	[L]
S_O	passive storage in old water reservoir	SV	[L]
S_S	storage in slow reservoir	SV	[L]
S_U	storage in unsaturated reservoir	SV	[L]
$S_{U_{max}}$	storage capacity in unsaturated reservoir	Р	[L]
T_F	<u>concentration time</u> for fast reservoir	Р	[T]
T_S	<u>concentration time</u> for slow reservoir	Р	[T]
X_F	reservoir creating lag time between $S_U \to S_F$	SV	[L]
X_S	reservoir creating lag time between $S_U \to S_S$	SV	[L]
β	shape parameter	Р	[-]

Table B.2. Expression table for the Loch Ard-Burn model. It contains expressions for all thefluxes. It requires an ancillary table for all the new definitions included in the expressions.

Variable	Name	Expression
E_{SI}	evaporation from vegetation	$\min(S_I/dt, E_p)$
E_{SU}	transpiration from unsaturated reservoir	$E_p C_E \min(1, S_U / (S_{U_{max}} L_p))$
E_{SF}	transpiration from fast responding reservoir	$\min(E_p(1-C_E), S_F/dt)$
P_R	Rainfall	•
P_{TF}	throughfall	$P_R - \min((I_{max} - S_I)/dt)$
Q_{SF}	runoff from fast reservoir	$K_F S_F$
Q_{OF}	overland flow	$max(S_F - S_{F_{max}}, 0)$
Q_{SS}	runoff from slow reservoir	$K_S S_S$
R_F	recharge of fast reservoir	$C_R(1-C_p)P_{TF}$
R_F^*	delayed flux from fast reservoir	$R_F \star h_F$
R_O	flux from hidden old water reservoir to unsaturated zone	$\equiv R_S$
R_P	preferential recharge of slow reservoir	$C_R C_P C_E$
R_S	recharge of <u>old water</u> reservoir	$P_{max}(S_U/S_{U_{max}})$
R_S^*	delayed flux from slow reservoir	$R_S \star h_S$
$\tilde{R_U}$	percolation from the unsaturated reservoir	$(1 - C_R)P_E$

Table B.3. Ancillary variables of the Loch Ard-Burn dictionary introduced by the flux expres-sions

Symbol	Name	Expression
$\begin{array}{c} C_R \\ h_F \\ h_S \end{array}$	coefficient of partition between runoff types response time distribution for X_f reservoir response time distribution for X_S reservoir	$(1 + \exp(-S_U/S_{U_{max}} + 0.5))^{-1}$ $2(t/T_F^2)$, for $0 < t < T_F$; 0 elsewhere $2(t/T_S^2)$, for $0 < t < T_S$; 0 elsewhere

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 Table B.4.
 Controllers table in the Ard Burn model

Symbol	Name	Expression
$\begin{array}{c} C_E \\ E_p \end{array}$	coefficient of partition between evapotranspirations potential evapotranspiration	$\frac{S_U/(S_U+S_F)}{?}$

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Snow routine (TT, CFMAX, SCF, CWH, CFR)

Soil routine (FC, LP, BETA)

Groundwater routine

(K0, K1, K2, UZL, PERC)



(b) Loch Ard – Burn 11



<u>(d) C_{M,dyn}</u>

Acce



place (reservoir) labeled with state variable

transition (flux) labeled with variable

controller labeled with *controlling variable*

generic **arc**



quantitites

measured quantities used for validation

stationary fluxes)

. 🗖

binding **arc**

controller-specific **arc**

symbol for measured

invisible place (subject to

splitter (it separates fluxes)



transition to place connection (*input to place*)



place to transition connection (output from place)





place affecting a controller. Meaning that T is a function of U_v



controller affecting a transition (flux). Meaning that parameters in E_T are a function of T

a splitter separating flux J into two contributions

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a system whose E_T is controlled (also) by 219 Aprilia Feophysical Union. All rights reserved.

Acce







Hidden Reservoir

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Shallow subsurface/Quick flow



${\it Groundwater/Slow\ reservoir}$

©2019 American Geophysical Union. All rights reserved. $Surface \ Waters$





 $\stackrel{(\!02019}{lmu}$ American Geophysical Union. All rights reserved. Above ground vegetation



Carbon Budget

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