

1 Black carbon aerosols at an urban site in North Africa (Kenitra, Morocco)

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8 **Abstract:**

9 Although black carbon is one of the most important atmospheric particulate components driving climate  
10 change and air quality, there is a lack of comprehensive air quality data from North African urban areas.  
11 Seven-wavelength aethalometer-collected data were used to provide first insights into the levels and  
12 patterns of equivalent black carbon (eBC) mass concentrations, particularly into the seasonal and diurnal  
13 cycles, that exist over an urban area in the North of Africa, Kenitra city, Morocco. The average BC mass  
14 concentration was observed as  $0.90 \pm 0.80 \mu\text{g}/\text{m}^3$ . This BC value is overall largely lower than reported  
15 African and European urban location levels, although the role of COVID-19 restrictions could have  
16 influenced this result. The diurnal pattern revealed two BC peaks corresponding to morning and evening  
17 traffic rush hours. This variation is found to be seasonally dependent, with the maximum daily amplitude  
18 occurring in winter. The BC bimodal behaviour is observed in all the seasons but with varying degrees  
19 of the magnitude of the peaks (winter>autumn>summer). This observed seasonal heterogeneity was  
20 explained by varying meteorological conditions (wind speed, wind direction, relative humidity, air  
21 pressure, air temperature, rainfall, and boundary layer height), and air mass trajectory height and  
22 pathways. It was pointed out that the daily and seasonal variations are significantly influenced by local  
23 traffic sources, wind speed, and boundary layer dynamics. Hysplit backward trajectories revealed that  
24 the long-range aerosols transport through both south-western (Atlantic region), and western (over  
25 populated areas) regions can as well contribute to BC concentrations, although to a different extent.  
26 Finally, a common thread that runs through seasonal, and diurnal variations were that typical low hourly  
27 BC concentrations (distinguished in the afternoon) do not exceed the value of  $0.5 \mu\text{g}/\text{m}^3$  representing a  
28 kind of urban background level.

29 **Keywords:** Carbonaceous aerosol; Black carbon; Aethalometer; Temporal variation; Morocco.

30

31 **Highlights:**

32 Real-time measurements of black carbon at a North African urban site, Kenitra city.

33 Low BC levels were observed in Kenitra compared with other African areas.

34 A well-defined diurnal and significant seasonal variation in BC concentration.

35 Local traffic activities assumed the major source of BC at the sampling site.

36 Long-range sources were found to be more prominent during winter.

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## 38 1. Introduction:

39 Carbonaceous aerosol is a major, sometimes dominant, constituent of atmospheric particulate matter  
40 (PM) in urban areas (Querol et al., 2013; Reche et al., 2012; Genga et al., 2017). It is primarily produced  
41 by three processes: wood combustion, fossil fuel combustion, and biogenic emissions. Carbonaceous  
42 aerosols are commonly divided into organic carbon (OC), and a refractory light-absorbing fraction  
43 known as elemental carbon (EC, when assessed using thermal-optical methods) or black carbon (BC,  
44 when quantified using optical methods). Black carbon in itself has received a lot of attention in recent  
45 decades because of its remarkable impact on regional-to-global climate change (Martinsson et al., 2017;  
46 Gatari et al., 2019; Pani et al., 2020). It is characterized as a by-product of incomplete combustion of  
47 fossil fuels in transportation engines, thermal power plants, and biomass in the residential sector, as well  
48 as smoke from agricultural and forest fires. Janssen et al. (2011) stated that combining BC measurements  
49 with other air quality indicators is advantageous since BC concentrations, as a specific indicator of local  
50 combustion sources, can reveal more about the health effects of aerosols than PM, which does not just  
51 come from burning. WHO global air quality guidelines stated that even though black carbon may not be  
52 the causal agent, it is considered as a useful additional air quality metric for assessing the health risks of  
53 primary combustion aerosols from traffic, which are not fully taken into account by PM<sub>2.5</sub> mass levels  
54 (WHO global air quality guidelines, 2021). In addition, there is more evidence of health effects linked  
55 to combustion sources than non-combustion sources (Segersson et al., 2017). In a recent report, the  
56 World Health Organization (WHO) reported pieces of evidence linking black carbon particles to  
57 cardiovascular health effects and premature mortality, for both short-term (24-hour) and long-term  
58 (yearly) exposures (WHO global air quality guidelines, 2021). Although insufficient data are available  
59 to provide recommendations for air quality guideline levels and interim targets, WHO warranted more  
60 actions to enhance further research on BC risks and approaches for mitigation, and outlined good  
61 practice statements for addressing BC in air quality monitoring networks (WHO global air quality  
62 guidelines, 2021).

63 BC temporal characteristics have long been a popular topic among researchers. Several investigations  
64 have looked at the peculiarities of BC mass concentration in different parts of Africa (Table 1), such as  
65 Tetouan (Morocco), Bou Ismaïl (Algeria), Mbour and Dakar (Senegal), Bamako (Mali), Cotonou  
66 (Benin), Yaounde (Cameroon), Nairobi and Nanyuki (Kenya), Kwadela (South Africa), Abidjan (Ivory  
67 Coast). For instance, in northern Morocco, Benchrif et al. (2018) reported that the carbonaceous particles  
68 account for 18% to 22% of the fine aerosol mass and range between 3.01 - 3.22  $\mu\text{g}/\text{m}^3$ , regarding the  
69 aerosol transport pathways and their loading region. In Kenya, Gatari et al. (2003) reported that the  
70 measured concentration of BC in urban Nairobi ( $2.37 \pm 0.3 \mu\text{g}/\text{m}^3$ ) was almost twice that of rural Meru  
71 ( $1.47 \pm 0.1 \mu\text{g}/\text{m}^3$ ). They claimed that local biomass burning was the primary source of BC at all  
72 measurement sites, although emissions from vehicles and waste burning in Nairobi may have enhanced  
73 the carbonaceous aerosols. Djossou et al. (2018) revealed mean BC concentration in Abidjan ranged  
74 from about  $2 \pm 1 \mu\text{g}/\text{m}^3$  near two-wheeled vehicles areas to approximately  $7 \pm 4 \mu\text{g}/\text{m}^3$  in four-wheeled  
75 vehicles site. Additionally, the characterization of BC mass concentration in a low-income settlement  
76 from Kwadela in South Africa by Xulu et al. (2020) was found to show seasonal variation, where the  
77 BC level was 2.9 times higher in winter ( $1.89 \pm 0.5 \mu\text{g}/\text{m}^3$ ) than in summer ( $0.66 \pm 0.2 \mu\text{g}/\text{m}^3$ ). However,  
78 those datasets were the first attempts to assess BC in Africa, and more long-term datasets with time  
79 resolution consistent with BC lifetime in the atmosphere are needed.

80 To date, multiple studies have shown that two key elements, namely source emissions and  
81 meteorological conditions, can influence the concentration levels and variations of BC, the role of long-  
82 transport air masses being also relevant. For instance, in urban environments, fossil fuel combustion  
83 mainly from the vehicle and industrial emissions is the major source of BC, while in rural areas it could  
84 be linked to biomass and biofuel combustion (Zotter et al., 2017). However, in many urban and suburban  
85 areas, biomass combustion, from forest fires (in summer) or domestic heating (in winter) may also  
86 contribute significantly to BC levels (Kalogridis et al., 2018). In Africa, the dominant sources of  
87 carbonaceous aerosols in urban settings are overall fossil fuel combustion and biomass burning. In urban  
88 Kenya, Gatari et al. (2003) have identified the local biomass burning, while in the coastal towns of  
89 southern West Africa, fossil fuels come from traffic, waste burning at landfills, and smoking activities

90 were suggested as BC sources by Djossou et al. (2018). In the Niger delta of Nigeria, vehicle emissions  
91 in cities, bush burning in rural areas, and destruction of illicit local oil refineries, as well as gas flaring  
92 as part of oil processing, are all major sources of BC (Giwa et al., 2014; Oluleye and Folorunsho, 2020).  
93 As result, it could be said that the sources of BC could present significant regional dependence (Cui et  
94 al., 2021). The relationships between BC concentration and meteorological factors have also been  
95 assessed in many cities such as Mbour (Senegal), Accra (Ghana), Cotonou (Benin), Abidjan (Ivory  
96 Coast), Kwadela (South Africa) (Rivellini et al., 2017; Djossou et al., 2018; Xulu et al., 2020; Alli et al.,  
97 2021). The authors of these studies indicated that meteorological parameters including wind speed,  
98 temperature, relative humidity, and boundary layer height have an impact on BC concentrations in urban  
99 environments and that these factors vary on a diurnal and seasonal timescale.

100 In this study, attempts were performed to understand the long-term concentration variability of  
101 equivalent black carbon (eBC) mass concentration with state-of-the-art technologies in a North African  
102 urban area, Kenitra in the West region of Morocco. Online measurements of BC were conducted  
103 between mid-July 2020 to mid-February 2021 to investigate their long-term evolution, seasonal  
104 variation, meteorological effects, and emission sources. Using data acquired by a 7-wavelength  
105 aethalometer over a lengthy period, the objectives of this study are i) assessment of eBC mass  
106 concentration variability (daily, monthly, and seasonally), ii) determination of levels, and sources of BC  
107 and major factors influencing its variability, iii) comparison of BC levels among urban African and  
108 European areas, iv) creation of the first real-time database of BC within a typical North-African city,  
109 Kenitra, as there is currently no dataset available. To the best of the authors' knowledge, no long-term  
110 investigation of carbonaceous aerosols in North Africa has been reported utilizing online measurement  
111 techniques.

## 112 **2. Materials and Methods**

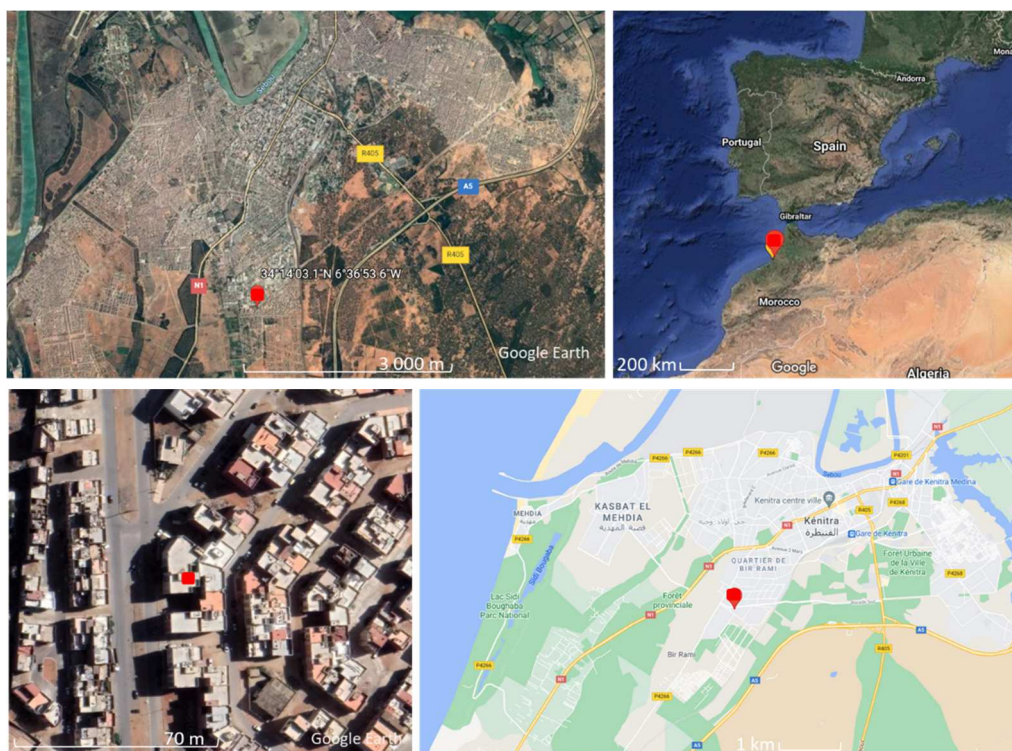
### 113 **2.1 Instrumentation and sampling**

114 A seven-wavelength (370, 470, 520, 590, 660, 880, and 950 nm) aethalometer (Model AE-31, Magee  
115 Scientific, USA) was used to perform online light-transmission measurements through a filter tape. The  
116 output of the aethalometer at each wavelength is expressed as equivalent black carbon mass  
117 concentration (BC) (Bernardoni et al., 2021). Measurements were carried out at a time resolution of 5-  
118 minutes from mid-July 2020 to mid-February 2021 in Kenitra, Morocco (Figure 1), in an urban area  
119 (Lat: 34.234194; Lon: -6.614889). Aethalometer was attached to PM<sub>2.5</sub> size-selective cyclone inlet (BGI,  
120 model SCC-1.828) pre-segregated all aerosols larger than 2.5  $\mu\text{m}$  from inflowing the instrument,  
121 requiring a flow rate of 5 L/min. The inlet was situated on a terraced roof containing the instrument, at  
122 a height of approximately 15 m above ground level. The Aethalometer was sampling under ambient  
123 conditions where drying of the aerosol stream was not performed. The lack of an active drying process  
124 led to operational relative humidity (RH) being somewhat lower than the ambient relative humidity due  
125 to slightly elevated instrument temperatures, especially during the night (Schmid et al., 2005). In our  
126 case, the average recorded temperature and relative humidity in the measuring room in summer were,  
127 respectively, 28 °C and 63 %. The calculated dew point temperature was within 20°C. Schmid et al.  
128 (2005) argued that the dependence of aethalometer-derived absorption coefficients on RH is not  
129 statistically significant, at least for RH between 40 and 80%.

130 The used aethalometer (AE-31) is based on the measurement of the optical properties of aerosols  
131 collected on a filter tape, over a short period. The aethalometer relies on measurements of transmission  
132 of light through the filter to obtain optical absorption coefficients ( $\sigma_{abs}$ , usually given as  $\text{M}\cdot\text{m}^{-1}$ ) at each  
133 wavelength and then equivalent black carbon (eBC) concentrations (Hansen, 2005). However, several  
134 studies (Wang et al, 2018; Collaud Coen et al., 2010) highlighted that aethalometer filter-based  
135 measurements exhibit three kinds of artifacts including filter loading, filter scattering, and scattering  
136 from the collected aerosols. The measurement biases caused by filter effects generate some uncertainty  
137 in BC measurements. Corrigan et al. (2006) reported an uncertainty ranging from only 5% up to 40% in  
138 absorption coefficients measured by an Aethalometer. Zanatta et al. (2016) proposed an uncertainty of

139 35% for absorption coefficients derived from the Aethalometer. Recently, Mousavi et al. (2018), in their  
140 study of BC using a seven-wavelength aethalometer, calculated the uncertainty in the BC concentration  
141 to be about 29%. Although several correction methods and algorithms have been established to correct  
142 for measurement biases, none of them are ideal and much additional information is required to use them.  
143 Backman et al. (2017) stated, for deriving an aethalometer correction factor, that the use of an additional  
144 collocated filter-based photometer measuring aerosol light absorption coefficients including Multi-  
145 Angle Absorption Photometer (MAAP) and Particulate Soot Absorption Photometer (PSAP). However,  
146 Wang et al. (2018) reported that different corrected results for the same original data may be led by  
147 various suggested algorithms.

148 In this study, we used the uncorrected AE-31 multi-wavelength absorption obtained data to estimate the  
149 BC contributions. Ganguly et al., (2005) estimated that the filter-loading correction due to multiple  
150 scattering effects could be ignored within eBC mass concentration in the range of 2 - 4  $\mu\text{g}/\text{m}^3$ . Ran et  
151 al. (2016) mentioned that measuring simultaneously and continuously aerosol absorption at different  
152 wavelengths has an important advantage. They suggested, to reduce systematic errors, handling properly  
153 multi-wavelength measurements that would be implied for exploring aerosol absorption spectral  
154 dependence, as well as inferring findings on chemical compositions and origins of light-absorbing  
155 aerosols.



156  
157 Figure 1. Location of the measurement site of the Carbonaceous Aerosol in Kenitra (red square) located  
158 in the Atlantic region (upper right panel), far from the city centre of Kenitra (upper left panel), between  
159 three major traffic roads (bottom right panel), and in an urban area (bottom left panel). Note the different  
160 scales of the maps.

## 161 2.2 Data processing and analysis

### 162 2.2.1 eBC Concentrations:

163 According to the instrument manufacturer, the Equivalent Black Carbon (eBC) from aethalometer AE-  
164 31 was obtained from measurements at wavelength  $\lambda = 880 \text{ nm}$ . eBC is the equivalent mass  
165 concentration of black carbon (BC) reported by the instrument. The eBC mass concentrations are

166 obtained by dividing the measured optical absorption coefficient ( $\sigma_{abs}$ ) at each wavelength by a mass  
167 absorption cross-section (MAC,  $m^2/g$ ).  $\sigma_{abs}$  is defined as the cross-section of carbonaceous aerosol  
168 available to absorb light per volume in which that eBC exists (Knox et al., 2009). MAC is wavelength-  
169 dependent which describes the eBC absorption per mass and scales inversely with wavelength. The eBC  
170 mass concentration from online measurements at the wavelength  $\lambda$  was obtained as follows (Eq. (1)):

$$171 \quad \sigma_{abs}(M.m^{-1}) = eBC(\lambda) \times MAC(m^2/g) \quad (1)$$

172 In the AE-31, the MAC values are derived from  $MAC(m^2/g) = \frac{14625}{\lambda}$ , where  $\lambda$  is the wavelength in  
173 nm. By using the formula given by the manufacturer, the calculated value for the MAC at 880 nm is  
174  $16.62 m^2/g$ . The filter tape is automatically advanced when the maximum value of optical attenuation  
175 reaches a suggested threshold value of 80, so that readings can take place on a pristine spot. Although  
176 Petzold et al. (2013) advocated using the acronym eBC for reporting the black carbon mass obtained  
177 from the mass absorption cross-section, we use the abbreviation BC in this work as we are addressing  
178 black carbon in general and it is widely used.

179 The collected data during the mid-July 2020 to mid-February 2021 period are investigated in this study.  
180 Only collected data from the aethalometer between the 1st and 99th percentiles were retained and  
181 selected to calculate a valid aggregated value (hourly average) and subsequent calculations. Anomalous  
182 values including negative and occasional single extreme values were regarded as invalid, and less than  
183 2% of data for Kenitra were discarded from the analysis. All the further averages, daily and monthly,  
184 were based on the >85% of validated hourly averaged data. Statistical data analyses were performed,  
185 comprising graphical presentations, in the base packages of the R programming language (R Core Team,  
186 2020) and four of its additional libraries: openair (Carslaw and Ropkins, 2012), ggplot2 (Wickham,  
187 2016), lubridate (Grolemund and Wickham, 2011) and dplyr (Wickham, 2021).

## 188 **2.2.2 NO<sub>2</sub> and Meteorological data**

189 Air quality pollutants (NO and NO<sub>2</sub>) were monitored with 30-min time resolution at the same BC  
190 measurement site throughout the corresponding sampling period. Meteorological data records were  
191 obtained from the weather station Kenitra Airport (34.274°N, 6.569°W), about 6.3 km from our study  
192 site. The 6-hour meteorological data includes wind speed (WS), wind direction (WD), relative humidity  
193 (RH), air pressure (P), and air temperature (temp). Boundary layer height (BLH) was estimated using  
194 the HYSPLIT™ (Hybrid Single-Particle Lagrangian Integrated Trajectory model, Version 5.0) (Stein et  
195 al., 2015; Rolph et al., 2017) by running Meteorological Profile. The BLH data were retrieved from the  
196 GDAS (Global Data Assimilation System) one-degree archive with a resolution of 3 h, which has a time  
197 step of 3-hours and a horizontal grid spacing of 1° latitude by 1° longitude. The R package worldmet  
198 (Carslaw D., 2021) was used for accessing NOAA Integrated Surface Database (ISD) meteorological  
199 observations (NOAA, 2020). Half-hourly NO and NO<sub>2</sub> concentrations were continuously monitored  
200 onsite using a NO<sub>x</sub> analyzer (HORIBA APNA-370). These auxiliary parameters were mainly used in  
201 correlation studies.

## 202 **2.2.3 Backward Trajectory Analysis**

203 To assess long-range transport patterns and explore the potential source regions contributing to the BC  
204 at the receptor site, the HYSPLIT™ (Version 5.0) was applied for air mass backward trajectory analyses  
205 (Benchrif et al., 2018; Jain et al., 2017). Air mass trajectories were provided by the GDAS current 7  
206 days archive (<ftp://arlftp.arlhq.noaa.gov/pub/archives/gdas1/>, accessed on October 2021). The 48-hour  
207 backward trajectories at hourly intervals were plotted at an altitude of 500 m-agl throughout the  
208 observation period (four trajectories per day every 6 hours). The trajectory length is not related to the  
209 BC or aerosol lifetime, but to the reliability of the HYSPLIT model (Fossum et al., 2022). It can be  
210 limited to a certain duration since the associated HYSPLIT error in the trajectory calculation is  
211 proportional to 15-30% of the trajectory path length (Benchrif et al., 2018; Draxler and Rolph, 2007).

212 To characterize the regional sources of carbonaceous aerosol and the transport of air masses arriving at  
213 a sampling location, many studies (Botsa et al., 2021; Deng et al., 2020; Doumbia et al., 2012) used 120-  
214 hour runs for the backward trajectory analysis, while others (Sun et al., 2020a; Duc et al., 2020;  
215 Martinsson et al., 2017) calculated hourly backward trajectories for the past 48-72 hours. Different  
216 arrival heights (500, 800, and 1000 m-agl) of backward trajectories were tested and the results show that  
217 the transport pathways are not very sensitive to the chosen heights within the studied area. To group the  
218 back-trajectories into distinct transport patterns, the trajectory cluster analysis method based on the  
219 squared Euclidean distance metric was applied. Back-trajectory classifications based on cluster analysis  
220 allow much more detailed information about potential source regions affecting Kenitra city. It should  
221 be noted that we considered regional air masses as a proxy for long-range BC aerosol transport. In this  
222 study, the cluster analysis was performed on trajectories for the period of the experiment using a total  
223 of 812 back-trajectories. Seasonal variability is explored by clustering trajectories for each season  
224 individually.

### 225 **3. Results and discussion**

226 In the first section, we compare the average BC data measured in Kenitra city to similar data from other  
227 African and Worldwide cities with the purpose to assess the representativeness of the collected data with  
228 respect to the African continent, and to put our dataset into a broader perspective. Also, we briefly  
229 discuss the possible influence of COVID-19 restrictions. In the following section, we assess the  
230 meteorology of Kenitra during the studied period in comparison to the typical climatic conditions of  
231 Northern African coastal cities. Finally, we assess the temporal variability of BC in Kenitra, addressing  
232 the possible role of local sources and long-distance factors.

#### 233 **3.1. Overview of eBC mass concentration over Africa**

234 The purpose of this section is to compare the average BC data measured in Kenitra city to data from  
235 other African and Worldwide cities. This is summarized in Table 1. In the Kenitra area, BC  
236 concentration was calculated (from mid-July to mid-February) to be on average  $0.90 \pm 0.80 \mu\text{g}/\text{m}^3$ .  
237 During winter (December to mid-February), autumn (September to November), and summer (mid-July  
238 to August) seasons, the mean BC value was  $1.30 \pm 1.17 \mu\text{g}/\text{m}^3$ ,  $1.01 \pm 0.69 \mu\text{g}/\text{m}^3$  and  $0.75 \pm 0.58$   
239  $\mu\text{g}/\text{m}^3$ , respectively. It may be pointed out from Table 1 that the levels of BC concentration are similar  
240 to those obtained at some African sites. For instance, the BC value of Kenitra is comparable to the mean  
241 BC obtained by Gatari et al. (2003) at a rural burning biomass site in Nanyuki, Kenya ( $0.7 \pm 0.06 \mu\text{g}/\text{m}^3$ ),  
242 and observed by Xulu et al. (2020) during periods of winter biomass burning ( $1.9 \pm 0.6 \mu\text{g}/\text{m}^3$ ) and  
243 summer solid-fuel combustion ( $0.7 \pm 0.2 \mu\text{g}/\text{m}^3$ ) at residential urban areas in Kwadela, South Africa. It  
244 may be seen also that the BC values in Kenitra are in the same range as those obtained at Bou Ismaïl  
245 (Algeria) coastal urban site ( $0.41 - 2.85 \mu\text{g}/\text{m}^3$ ) during Sep. 2011- Jan. 2012 period (Khedidji et al.,  
246 2020). Nonetheless, the BC concentration values of our site are largely lower to values reported in the  
247 literature at other urban traffic locations in Morocco (Tetouan) ranging between  $3.01 - 3.22 \mu\text{g}/\text{m}^3$  as  
248 reported by Benchrif et al. (2018), and Africa namely: Abidjan ( $7.0 \pm 2.6 \mu\text{g}/\text{m}^3$ ) (Adon et al., 2020),  
249 Bamako ( $19.2 \pm 8.9 \mu\text{g}/\text{m}^3$ ), Cotonou ( $4.9 \pm 3.9 \mu\text{g}/\text{m}^3$ ) (Doumbia et al., 2012), and Nairobi ( $2.3 \pm 0.3$   
250  $\mu\text{g}/\text{m}^3$ ) (Gatari et al., 2003). Average BC levels measured in this study are similar to the range of values  
251 reported in previous studies at urban background sites in Rome ( $1.0 \pm 0.6 \mu\text{g}/\text{m}^3$ ) (Costabile et al., 2015),  
252 Barcelona ( $1.7 \pm 0.6 \mu\text{g}/\text{m}^3$ ), Huelva ( $0.7 \pm 0.4 \mu\text{g}/\text{m}^3$ ), and North Kensington ( $1.9 \pm 0.7 \mu\text{g}/\text{m}^3$ ) (Reche  
253 et al., 2011). Wang et al. (2020) reported BC values of  $6.0 \pm 2.9 \mu\text{g}/\text{m}^3$ , from November to December  
254 2019, in Xi'an, China (Urban/ Residential site). Annual average BC measurements in urban and  
255 suburban sites in Shanghai, China quantified by Peng et al. (2019) are to be  $1.17 \pm 0.61 \mu\text{g}/\text{m}^3$  and  $2.09$   
256  $\pm 0.97 \mu\text{g}/\text{m}^3$ , respectively, attributed to the higher local combustion emissions and regional long-range  
257 transport. With this comparison, we can underline that Kenitra city (Morocco, North Africa) has  
258 particularly low carbonaceous pollution levels when compared to other traffic areas (e.g. West Africa)  
259 and megacities (e.g. Europe).

260 Furthermore, it may be noted that BC measurements were made during the COVID-19 curfew imposed  
 261 by the government starting on June 11, 2020, and remained in place until the end of the sampling. A  
 262 detailed timeline of the mobility restrictions and relaxation measures that have been taken in Kenitra  
 263 city during the Coronavirus time is briefed in Table S6. Accordingly, the low levels found in this study  
 264 are most likely linked to the nationwide curfew (between 7 p.m. and 5 a.m. or 9 p.m. to 6 a.m.) imposed  
 265 by authorities during the sampling period because of the COVID-19 circumstances. The curfew time  
 266 period in this study might not be truly representative of situations during typical business hours. During  
 267 this period, most of the activities took place during the daytime which affect Morocco as a whole and  
 268 reflect the unusual life of the city. Indeed, the lack of clear weekend/weekend effects would be expected  
 269 as a long night curfew period was established in Kenitra. Panda et al. (2021) pointed out that the  
 270 imposition of a night curfew influences the BC concentrations in Bhubaneswar, India, and declined by  
 271 ~47%. Kompalli et al., (2014a) argued that the impact of cut-off of sources owing to issues like night-  
 272 curfew or national strikes is dramatically remarkable. They found that BC concentrations in Indian semi-  
 273 urban areas dropped quickly to ~22% of the average concentrations that prevailed before a regional  
 274 “strike-work” that halted all industrial activities and commercial road traffic for two days. Otherwise,  
 275 following the COVID-19 lockdown stage, Xu et al. (2020) found that BC values in the Chinese city  
 276 curbed down by 44%, from 2.30 to 1.29  $\mu\text{g}/\text{m}^3$ . Gogoi et al. (2021) as well reported that BC levels from  
 277 pre-lockdown observations are reduced significantly (10%-40%) in several Indian locations, mainly  
 278 over the regions where there are abundant non-transport fossil fuel emission sources (e.g. industries,  
 279 thermal power plants) compared to the those where the transport sector makes a major contributor.

280 Table 1: Comparison between BC values measured in Kenitra and across African and worldwide cities,  
 281 focusing on urban sites described as being impacted by traffic. The table covers the type of sites, study  
 282 period, measurement techniques/methods, and the most noteworthy findings. The table is ordered  
 283 according to BC mass concentration.

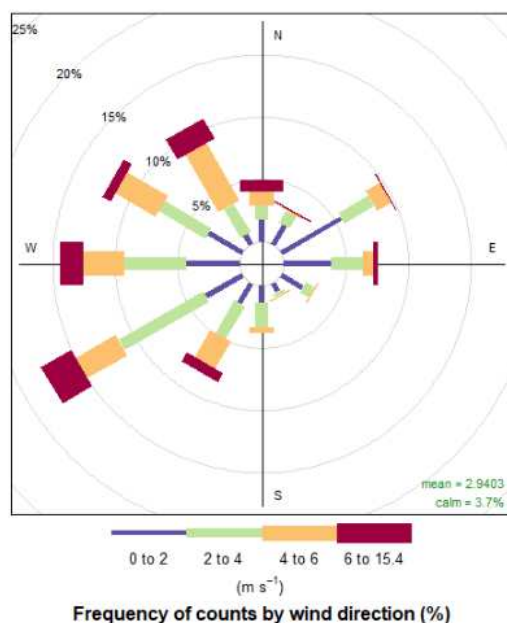
Location	Type of location	Study period	Measurement techniques/methods	BC mass concentration range, $\mu\text{g}/\text{m}^3$	Reference
Middle Atlas (AM5), Morocco	Remote high-altitude	Aug.- Dec. 2017	Thermal-optical transmittance technique (Sunset laboratory, USA) in PM10 samples.	$0.17 \pm 0.10$	Deabji et al. (2021)
Mbour, Senegal	Coastal site located in the outskirts of the city	Mar.- Jun. 2015	7-wavelength Aethalometer (AE-33) equipped with a PM1 impactor inlet (BGI model SCC-0.732) and sampled at 5 L/min.	$0.36 \pm 0.37$	Rivellini et al. (2017)
Nanyuki, Kenya	Rural	May-June 1999	A black smoke detector model FH 62 I-N (ESM Emberline, Erlangen, Germany) for measuring fine BC.	$0.7 \pm 0.06$ (0.4-1.1)	Gatari et al. (2003)
Huelva, Spain	Urban background site influenced by industrial emissions	2009	Multi-angle absorption photometer (MAAP Thermo ESM Andersen Instrument) with PM10 inlet.	$0.7 \pm 0.4$	Reche et al. (2011)
<b>Kenitra, Morocco</b>	Urban	July 2020 to February 2021	7-wavelength aethalometer (AE-31) with a PM2.5 inlet and a flow rate of 5 L/min.	$0.90 \pm 0.80$ (winter: $1.30 \pm 1.17$ ; autumn: $1.01 \pm 0.69$ and summer: $0.75 \pm 0.58$ )	This work
Rome, Italy	Suburban background	Dec. 2011 – Feb. 2012	A 3-wavelength particle soot absorption photometer (PSAP) and Nephelometer (Ecotech, mod.Aurora 3000) operated to measure the dry aerosol scattering coefficient at three wavelengths (450,520, 635 nm).	$1.0 \pm 0.6$	Costabile et al. (2015)
Shanghai, China	Urban	Jun. 2016- Jun. 2017	7-wavelength aethalometer model AE31 from Magee Scientific with inlet cut-off size of 2.5 $\mu\text{m}$ and operated at a flow rate of 5 L/min.	$1.17 \pm 0.61$	Peng et al., 2019
Bou Ismaïl, Algeria	Coastal area	Sep. 2011 to Jan. 2012	Thermal-optical transmittance technique (Sunset laboratory, USA) in PM10 samples.	$1.25$ (0.41 - 2.85)	Khedidji et al. (2020)
Cotonou, Benin	Urban traffic	Feb. 2015-Mar. 2017	IMPROVE_A Temperature Protocol for Thermal/Optical Carbon Analysis (in PM2.5 aerosols).	$1.5 \pm 0.2$	Djossou et al. (2018)
Barcelona, Spain	Urban background site in a city with dense traffic	2009	Multi-angle absorption photometer (MAAP Thermo ESM Andersen Instrument) with PM10 inlet.	$1.7 \pm 0.6$	Reche et al. (2011)
Kwadela, South Africa	Semi-Urban	Winter 2014 (Jul.- Sep.) and Summer 2015 (Feb.- Apr.)	Multiple wavelengths aethalometer (AE-22 and AE-31) with a PM2.5 inlet.	$1.9 \pm 0.6$ (winter, AE-22) $0.7 \pm 0.2$ (summer, AE-31)	Xulu et al. (2020)
North Kensington, UK	Urban Background	2009	Magee AE-21 Aethalometer	$1.9 \pm 0.7$	Reche et al. (2011)
Shanghai, China	Suburban	Jun. 2016- Jun. 2017	7-wavelength aethalometer model AE31 from Magee Scientific with inlet cut-off size of 2.5 $\mu\text{m}$ and operated at a flow rate of 5 L/min.	$2.09 \pm 0.97$	Peng et al., 2019

Hangzhou, China	Urban, Urban industry, Suburban, and Rural	1 Jan. - 31 Mar. 2020 (during, pre- and post-COVID-19 lockdown)	7-wavelength aethalometer (AE-31) with a flow rate of 5 L/min.	The citywide BC concentrations decrease from 2.30 to 1.29 $\mu\text{g}/\text{m}^3$	Xu et al. (2020)
Nairobi, Kenya	Urban	March 2000	A black smoke detector model FH 62 I-N (ESM Emberline, Erlangen, Germany).	$2.3 \pm 0.3$	Gatari et al. (2003)
Rome, Italy	Urban background	February 2017	7-wavelength aethalometer (AE-33) and Multi-Angle Absorption Photometer (MAAP).	$2.6 \pm 2.5$	Costabile et al. (2017)
Tetouan, Morocco	Urban	May 2011 – Apr. 2012	A two-step combustion approach in PM2.5 developed by Cachier et al. (1989).	3.01-3.22	Benchrif et al. (2018)
Cotonou, Benin	Urban traffic	May 2005	7-wavelength aethalometer (AE-42) with a PM2.5 inlet and operated at a flow rate of 4 L/min.	$4.9 \pm 3.9$	Doumbia et al. (2012)
Yaounde, Cameroon	Urban traffic	Aug. 2010	7-wavelength aethalometer (AE-31) with a PM2.5 inlet and a flow rate of 5 L/min.	$5.7 \pm 2.3$	Doumbia et al. (2012)
Dakar, Senegal	Urban traffic	Jun. 08- Jul. 09	7-wavelength aethalometer (AE-42) with PM2.5 inlet and operated at a flow rate of 2 L/min.	5.7-15.4	Doumbia et al. (2012)
Xi'an, China	Urban/ Residential	1 Nov. - 31 Dec. 2019	7-wavelength aethalometer (AE-33) with PM2.5 inlet.	$6.0 \pm 2.9$	Wang et al. (2020)
Abidjan, Ivory Coast	Urban traffic	Jan. 2016	Two-step thermal method in PM2.5 developed by Cachier et al. (1989).	$7.0 \pm 2.6$	Adon et al., (2020)
Abidjan, Ivory Coast	Urban traffic	Feb. 2015-Mar. 2017	IMPROVE_A Temperature Protocol for Thermal/Optical Carbon Analysis (in PM2.5 aerosols).	$7 \pm 4$	Djossou et al. (2018)
Addis Ababa, Ethiopia	Urban/ Residential	Nov.2015- Nov. 2016	Thermal-optical transmittance technique (Sunset laboratory, Forest Grove, OR, USA) in PM2.5 samples.	$13.7 \pm 6.8$	Tefera et al., (2020)
Bamako, Mali	Urban traffic	April 2008	7-wavelength aethalometer (AE-42) with a PM2.5 inlet and operated at a flow rate of 4 L/min.	$19.2 \pm 8.9$	Doumbia et al. (2012)

### 284 3.2 Meteorological situation

285 The meteorology of Kenitra (Table 2) is characterized by high temperatures in summer (mid-July to  
286 August), with monthly values ranging 37.1-13.0 °C, against 21.1-2.7 °C in winter (December to mid-  
287 February). Relative humidity stands between 75 % in summer, up to 79 % in autumn (September to  
288 November), and winter. Kenitra is usually windy, with wind speed averaging 2.7 m/s in autumn and  
289 winter, while it represents 3.3 m/s on average during summer. This is typical of North African cities  
290 which have semiarid and arid desert climates characterized by extremes in daily high and low  
291 temperatures, hot summers and cold winters, and minimal rainfall of approximately 200 to 400 mm per  
292 year for semi-arid regions and less than 100 mm per year for desert regions (NIC special report, 2009).

293 The time-series variations of temperature, rainfall, wind speed, air pressure, relative humidity, and  
294 boundary layer height derived from hourly measurements for the whole sampling period are shown in  
295 Figures S1a and S1b. The temperature variations show a decreasing trend from autumn to winter while  
296 the wind speed displays an increase including several episodes of stronger winds predominantly during  
297 the transition period (December-January). The seasonal cycle reveals that winter is associated with more  
298 stable atmospheric conditions due to lower wind speed compared to summer. Many episodes of rainfall  
299 and high relative humidity (>79 %) were very frequent during autumn and winter, particularly from  
300 December to January. The prevalent wind direction was the Southwest during the measurement  
301 campaign, while the Northwest prevailing wind direction was also registered. Higher frequencies of  
302 stronger winds (4–6 m/s and 6–8 m/s) were recorded during the autumn period. The wind rose (Figures  
303 2 and S3) shows clearly that measurements were influenced dominantly by a fairly weak wind flow of  
304 2.9 m/s on average from predominant west-southwestern directions. Winds from other directions are  
305 rarer. The air pressure ranged from 1002 to 1030 hPa, exhibiting an increase from summer to winter. A  
306 decreasing tendency was observed from mid-December to mid-January, with fluctuations in autumn.  
307 BLH ranged from 933.3 m in August to 953.1 m in December, with an annual average of  $956.8 \pm 165.3$   
308 m.



309

310

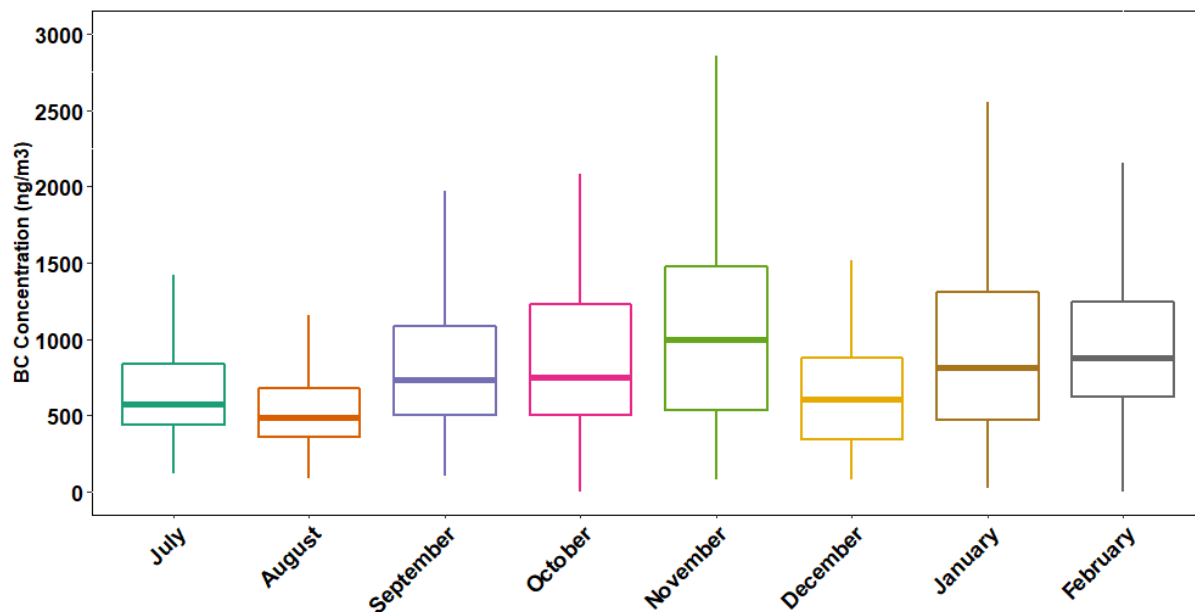
Figure 2: Wind rose based on hourly averages of wind speed and wind direction.

311 **3.3 Seasonal features of BC concentrations at Kenitra**

312 This section summarizes the BC concentration patterns from continuous measurements for eight months  
 313 at Kenitra city. The corresponding monthly profile is shown in Figure 3 and descriptive statistics can be  
 314 found in Table 2. As shown in Figure 3 and Table 2, the BC concentrations displayed a remarkable  
 315 seasonal pattern, as their levels were high during winter and low during summer. The average ( $\pm$ SD,  
 316 standard deviation) of BC concentrations were, respectively,  $1298.1 \pm 1169.3 \text{ ng/m}^3$  in winter (45%  
 317 higher than the entire period average of  $896.9 \pm 800.3 \text{ ng/m}^3$ ), against  $1006.0 \pm 688.4 \text{ ng/m}^3$  (12%  
 318 higher) in autumn and  $752.1 \pm 575.5 \text{ ng/m}^3$  (16% lower) during summer. The shape and trend of the  
 319 seasonal variation of BC concentration are consistent with findings of Xulu et al. (2020) in Kwadela  
 320 Township in South Africa and Doumbia et al. (2012) over Dakar city in Senegal. Several hypotheses are  
 321 discussed below to explain the significant seasonal and monthly variation in BC concentrations.

322 Cao et al. (2009) argued that wind speed, wind direction, and temperature were important meteorological  
 323 factors that affected BC concentrations. In our case study, however, a weak positive and negative  
 324 association was revealed between BC and meteorological parameters, and the BC concentration  
 325 variation was likely less affected by changes in this prevailing meteorology. Thus, our data were  
 326 separated into three parts according to the season and conducted a correlation analysis in each of these  
 327 seasons. Tables S2, S3, S4, and S5 provide a summary of the correlation matrices for the whole study  
 328 period, and during the different seasons. The normality of the studied variables was checked using the  
 329 Shapiro-Wilk test, which showed all variables don't follow the normal distribution. Then, Spearman  
 330 correlations were employed. The calculated Spearman's correlation coefficients indicated that  
 331 carbonaceous fraction concentrations are significantly correlated with meteorological factors, at 95% or  
 332 99% confidence levels. As shown, the relation between BC and WS, RH, T, and P varies with the  
 333 seasons, and the correlation in winter is slightly stronger than in summer and autumn. For instance, a  
 334 weak negative correlation between BC concentrations and wind speed was found ( $r$  ranged from -0.46  
 335 ( $p < 0.01$ ) in winter to -0.31 ( $p < 0.01$ ) in autumn and -0.23 ( $p < 0.01$ ) in summer). These results support the  
 336 fact that seasonal variations do exist in the relationships between BC concentrations and meteorological  
 337 factors. In addition, the BC concentration throughout winter was likely driven primarily by the low wind  
 338 speed, confirming the key role of wind speed, among other meteorological factors in the observed BC  
 339 concentrations (Sun et al, 2020b). The seasonal variation in BC mass concentration could also be mainly  
 340 related to seasonal differences in BLH as well as meteorological factors. The observed stable BLH, due  
 341 to the low wind speed and hence low surface convection, prevents aerosol particles from dispersing

342 strongly into the atmosphere and is enclosed near the boundary layer showing the high BC concentration  
 343 in winter (Guha et al., 2014). In general, the atmosphere of Kenitra during the winter season is  
 344 characterized by high relative humidity and low ventilation coefficients, leading to lower aerosol  
 345 dispersion and, as a result, an increase in BC concentrations. However, the decrease in BC from winter  
 346 to summer can be related to the combined effects including the increase in ventilation coefficient (the  
 347 product of BLH height and wind speed) in the hot season (Beegum et al., 2009; Kompalli et al., 2014b).  
 348 During the summer season, the enhancement of both BLH and ventilation coefficient are conducive  
 349 conditions for lofting on near-surface aerosols to higher atmospheric levels. Aside from these former,  
 350 the local emissions and long-distance aerosol transport also contribute to the seasonal differences, which  
 351 are discussed in subsequent paragraphs.



352  
 353 Figure 3: Monthly profile of BC measured at 880 nm in Kenitra city between July 2020 and February  
 354 2021. Box-plot indicates minimum, maximum, median (dash line), and 25/75% quartiles (boxes) in the  
 355 dataset. Descriptive statistics of monthly BC concentrations are summarized in Table S1. It is worth  
 356 noting that because of the failure of the instrument system, a remarkable decline in BC levels is observed  
 357 in December due to the reduction in the number of valid observations as shown in Table S1.

358 Correlation analysis of BC with NO, NO<sub>2</sub>, and NO<sub>x</sub> (Figure S4) shows that the relationship between BC  
 359 and NO<sub>2</sub> followed two different patterns. The mass concentrations of BC were good correlated with the  
 360 NO<sub>2</sub> concentrations during the cold season as on November 2020 ( $r=0.67$ ,  $p\text{-value} < 5.194e-11$  which is  
 361 less than the significance level  $\alpha=0.05$ ) and January 2021, ( $r=0.57$ ,  $p\text{-value} < 2.2e-16$ ), but a weaker  
 362 correlation ( $r=0.49$ ,  $p\text{-value} < 1.851e-09$ ) was found for the other sampling days (August 2020). It can  
 363 be inferred that these are the results of the increase in traffic-related pollutants in the winter. Increased  
 364 traffic activities can result in increased emissions of NO<sub>2</sub> and other air pollutants. Accordingly, we can  
 365 speculate that the increase in NO<sub>2</sub> levels makes the correlation between BC concentrations and some  
 366 meteorological factors stronger.

367 Another hypothesis involves that the observed BC variations are probably due to the changes in regional  
 368 and long-range emission sources as discussed by Doumbia et al. (2012) and Pani et al. (2020).  
 369 Accordingly and to identify different source regions that are likely to impact BC levels in Kenitra city,  
 370 the cluster analysis technique based on the squared Euclidean distance metric was applied to HYSPLIT  
 371 back-trajectories of air masses arriving at the sampling site. Figure 4a shows mean back-trajectories  
 372 arriving at Kenitra from mid-July 2020 to mid-February 2021 classified into 6 backward trajectory  
 373 clusters. The major transport patterns can be summarized as follow: Slow North-easterly (24.0% of the

374 time, cluster (or C) 1), North-westerly (57.1%, C2, 3, and 4), and South-westerly (15.5%, C5 and 6). As  
375 shown in Figure 4b, the air masses from the Western coast of Morocco (C6) remain below 800 m for 48  
376 hours before reaching Kenitra. The air masses from northern regions of Morocco (C1) travel at altitudes  
377 between 600 and 800 m for the last 24 h until they approach Kenitra. On the contrary, the air masses  
378 from the Atlantic Ocean (C5 and C4) travel above 800 m in altitude most of the time. The clusters (C5  
379 and C6) display different transport elevation profiles although air masses followed the same path the  
380 day before reaching the sampling site.

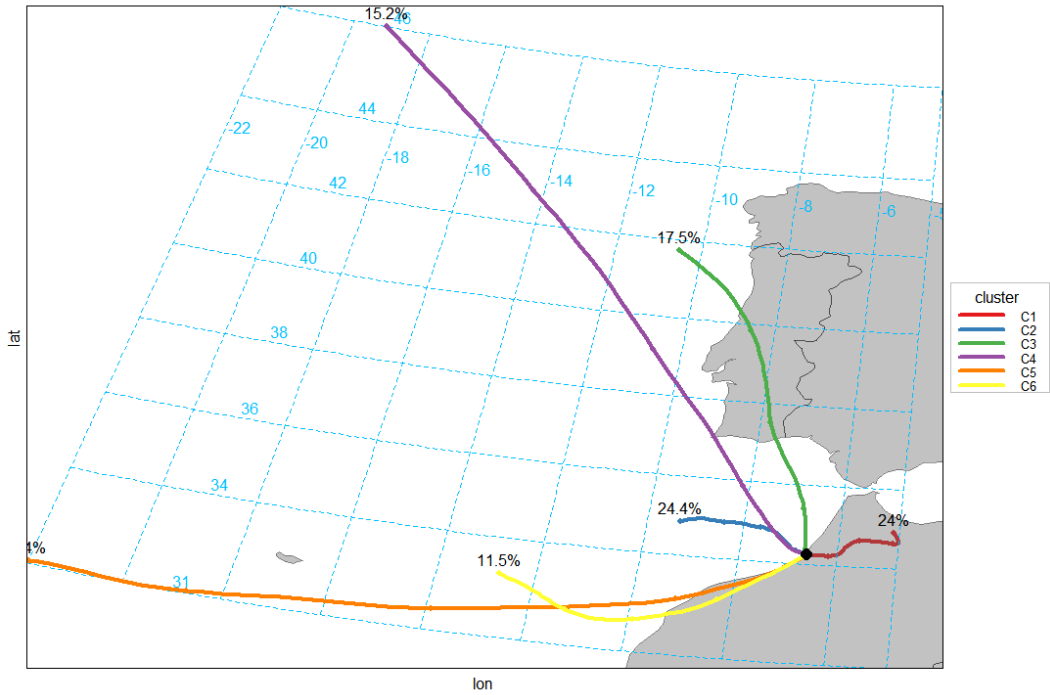
381 Several studies (Lack et al., 2012 and Mamoudou et al., 2018) stated that BC emissions from  
382 international ships are very significant. Jonson et al. (2020) reported that in addition to NO<sub>2</sub> and sulfur  
383 considered as the main sources of particles from shipping, it is assumed that BC is also emitted as  
384 primary particles. Cui et al. (2021) confirmed that regional transports have a significant impact on BC  
385 levels. They found that long-range aerosol transport, especially when air masses were linked to regions  
386 with high ship emissions as in the Yellow Sea and the Bohai Sea, could not be negligible for contributing  
387 to BC pollution in rural Qingdao, north-eastern China. A key characteristic of backward trajectory  
388 clusters classification is that the contribution of long-range carried aerosols outside areas is combined  
389 with primary emissions of carbonaceous components released by the maritime activity in the Atlantic  
390 region. These long-range contributions, however, are far from large enough to explain the observed  
391 seasonal variations in BC concentrations, as BC values are assumed to be the result of the contribution  
392 from the different source sectors, namely long-range and local emissions.

393 Figure 4c illustrates the back-trajectory analyses throughout different seasons of the sampling period.  
394 Contrary to the other seasons, the clustering of back-trajectories during winter led to distinct clusters  
395 (C1 and C4) corresponding to air masses within 500 m and relatively high concentrations of  
396 carbonaceous aerosols (Table 3). BC concentrations from C1 to C4 have comparable values of around  
397 1.4 µg/m<sup>3</sup>. A high BC value (1.4 µg/m<sup>3</sup>) was measured in autumn, driven by the air masses from the  
398 West coast of Morocco (C3). The other autumn clusters, C1, C2, C4, and C5, also showed high BC  
399 levels, averaging 1.0-1.1 µg/m<sup>3</sup>. It is, therefore, reasonable to assume that the polluted air masses, during  
400 the last 24 h before reaching Kenitra, are likely to be a key feature affecting the BC concentrations and  
401 causing part of the pronounced seasonal variation are: i) slowly traveling from the West coast of  
402 Morocco over the populated areas of Casablanca towards Kenitra; and ii) coming from the near Atlantic  
403 Ocean, affected by pollutants emitted from Northern cities of Morocco and Gibraltar Strait. In  
404 conclusion, until this stage, the observed trend whereby BC concentration increases during winter and  
405 drops during summer can be explained by varying emission sources, meteorological conditions, and air  
406 masse trajectory height and pathways.

407 Using bivariate polar plots, by calculating the mean concentration for wind speed and directions bins, it  
408 was possible to identify and understand the sources of BC coming from different directions (Grange et  
409 al., 2016). These graphs show how the concentration of BC varies according to the speed and direction  
410 of the wind. In their simplest form, BC concentrations are aggregated into wind speed and wind direction  
411 bins and are displayed on polar coordinates. Figure 5 presents polar plots of BC concentration for the  
412 whole period (upper panel) and in three seasons (winter, summer, and autumn) at Kenitra city. It  
413 suggests that locally-sourced carbonaceous aerosol was present, as demonstrated by the elevated  
414 concentrations at low wind speeds, while the low concentrations were observed with westerly winds at  
415 high wind speeds (>8 m/s). In the winter, it was observed that high concentrations of BC come from  
416 everywhere, being more relevant from the east with weak winds with speeds ranging from 0 to 4 m/s.  
417 In the autumn, high BC concentrations come from the northwest associated with strong winds with  
418 speeds ranging from 3 to 7 m/s. By contrast, summer showed that only when wind speeds were high,  
419 moderate to low concentrations were experienced due to enough degree of pollutant dispersion.  
420 Concisely, polar plots indicated that the major source of BC is in the vicinity of the monitoring site,  
421 consistent with road traffic. Moreover, long-range south-westerly atmospheric transport events can  
422 contribute to elevated BC concentrations.

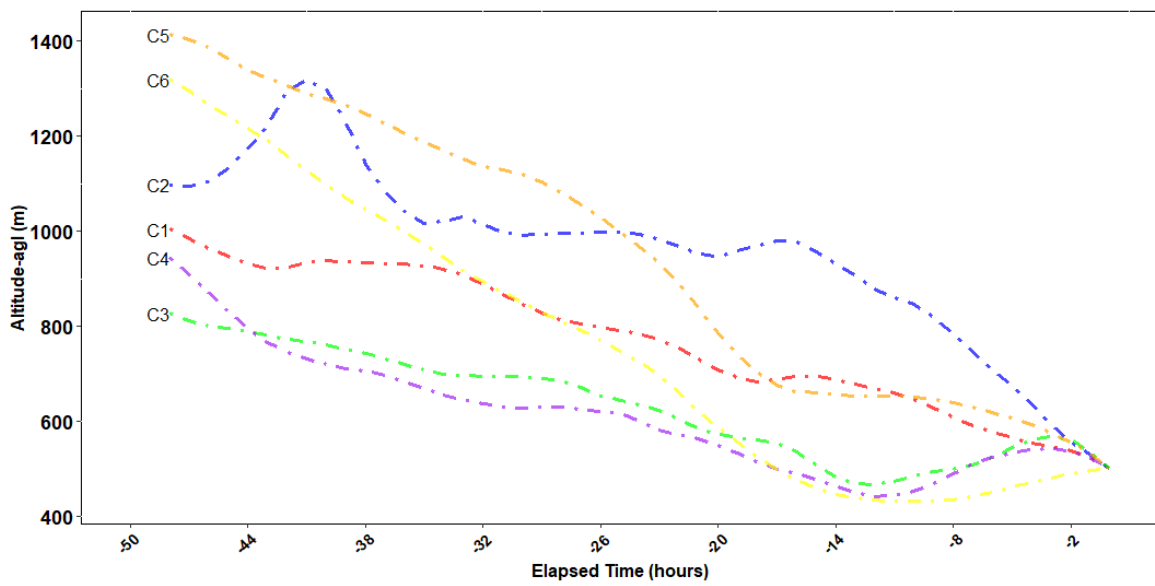
423 Table 2: Summary of the mean ( $\pm$  standard deviation), minimum (min), and maximum (max)  
 424 concentrations of 5-minutes BC (i.e., total BC), Daily-Ratio (Ratio of hourly BC between night-time  
 425 and daytime), and 6-hours meteorological parameters (T (temperature), P (pressure), BLH (Boundary  
 426 layer height), RH (relative humidity), WD (wind direction) and WS (wind speed)) between mid-July  
 427 2020 to mid-February 2021 (203 days). Summary data are also subdivided into seasons over Kenitra  
 428 city. n.a = not available. N stands for the number of total datasets used, where N = 11434 for BC; and  
 429 N = 781 for meteorological parameters.

Variable	Entire period	Winter	Autumn	Summer
	Mean $\pm$ SD [Max - Min]			
N for BC	11434	2730	649	8055
BC (ng/m <sup>3</sup> )	896.9 $\pm$ 800.3 [6693.0 - 0.7]	1298.1 $\pm$ 1169.3 [6693.0 - 0.7]	1006.0 $\pm$ 688.4 [4214.0 - 1.9]	752.1 $\pm$ 575.5 [6494.0 - 89.0]
Daily-Ratio of hourly BC	1.19 [1.37 - 0.74]	1.49 [1.78 - 1.07]	1.32 [1.80 - 0.88]	1.02 [1.29 - 0.52]
N for meteorological parameters	781.0	144	325	312
T (°C)	18.3 $\pm$ 6.0 [37.1 - 2.7]	12.4 $\pm$ 4.0 [21.2 - 2.7]	16.2 $\pm$ 4.6 [29.5 - 4.0]	23.3 $\pm$ 3.9 [37.1 - 13.0]
RH (%)	79.2 $\pm$ 17.0 [100.0 - 0.0]	86.8 $\pm$ 16.1 [100.0 - 0.0]	79.0 $\pm$ 19.0 [100.0 - 0.0]	75.8 $\pm$ 13.8 [100.0 - 0.0]
BLH (m-agl)	956.8 $\pm$ 165.3 [1396.0 - 118.0]	946.8 $\pm$ 183.3 [1396.0 - 119.0]	960.9 $\pm$ 171.7 [1382.0 - 118.0]	956.1 $\pm$ 148.9 [1347.0 - 465.0]
WS (m/s)	2.9 $\pm$ 2.1 [15.4 - 0.0]	2.7 $\pm$ 2.1 [9.8 - 0.0]	2.7 $\pm$ 2.1 [15.4 - 0.0]	3.3 $\pm$ 1.9 [9.8 - 0.0]
WD (degree)	218.4 $\pm$ 100.4 [360.0 - 10.0]	168.5 $\pm$ 97.2 [360.0 - 10.0]	199.6 $\pm$ 105.6 [360.0 - 10.0]	261.7 $\pm$ 76.6 [360.0 - 10.0]
P (hPa)	1017.2 $\pm$ 5.03 [1030.5 - 1002.2]	1019.7 $\pm$ 6.9 [1030.5 - 1002.2]	1018.7 $\pm$ 4.3 [1027.7 - 1006.7]	1014.3 $\pm$ 2.9 [1023.4 - 1007.8]
Sunrise time (a.m, local time)	n.a	08:27 [08:33 - 08:15]	07:34 [08:14 - 07:00]	06:42 [07:00 - 06:26]
Sunset time (p.m, local time)	n.a	06:33 [07:04 - 06:15]	06:53 [07:50 - 06:15]	08:17 [08:37 - 07:52]



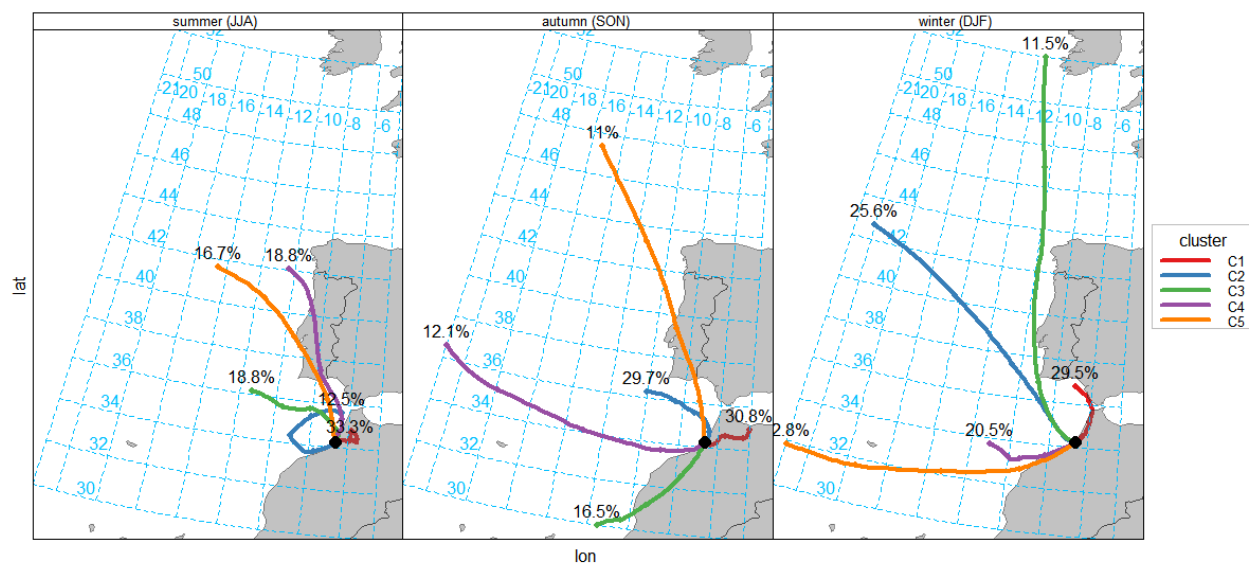
430

431 Figure 4a: mean back-trajectories arriving at Kenitra from mid-July 2020 to mid-February 2021  
 432 classified into 6 backward trajectory clusters.



433

434 Figure 4b: Mean altitudes of each back-trajectory cluster during the whole period.



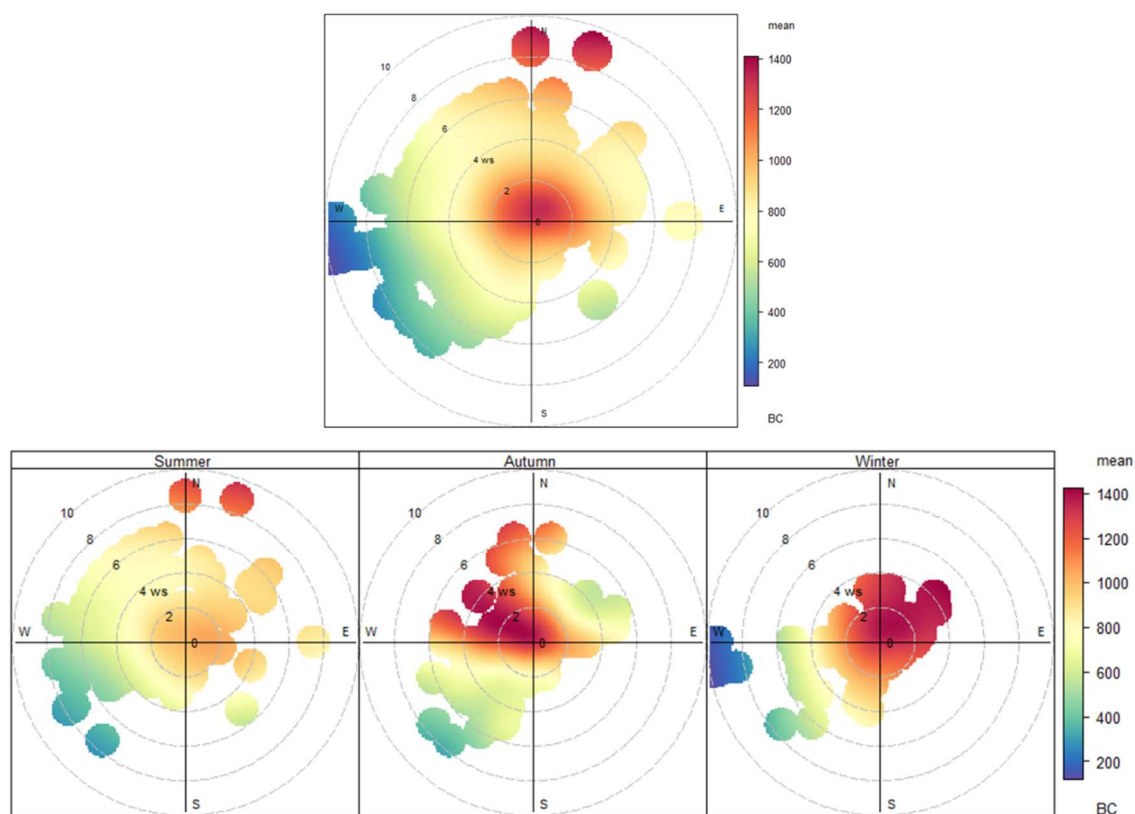
435

436 Figure 4c: Mean back-trajectories arriving at Kenitra from mid-July 2020 to mid-February 2021  
 437 classified into 6 backward trajectory clusters and subdivided into seasons: summer (from mid-July to  
 438 August), winter (from December to mid-February), and autumn (from September to November).

439 Table 3: Summary of the mean ( $\pm$  standard deviation), minimum (min), and maximum (max)  
 440 concentrations of BC related to the obtained clusters during the whole period and the three seasons. n.a  
 441 = not available.

Variable	Entire period	Winter	Autumn	Summer
BC (ng/m <sup>3</sup> )	Mean $\pm$ SD [Max - Min]			
N of Clusters	6	5	5	5
C1	989.2 $\pm$ 561.0 [2934.1 - 381.3]	1444.5 $\pm$ 958.7 [3100.8 - 554.0]	1035.4 $\pm$ 463.2 [2565.3 - 381.3]	567.8 $\pm$ 154.6 [989.3 - 397.4]
C2	1015.2 $\pm$ 624.1 [3421.6 - 416.6]	940.1 $\pm$ 305.2 [1445.2 - 562.1]	1043.6 $\pm$ 274.1 [1812.2 - 720.4]	630.8 $\pm$ 146.3 [864.5 - 416.6]
C3	926.5 $\pm$ 535.0 [3100.8 - 170.2]	475.8 $\pm$ 341.5 [950.7 - 170.2]	1446.2 $\pm$ 682.5 [2665.0 - 500.2]	627.3 $\pm$ 186.4 [968.6 - 458.6]
C4	928.8 $\pm$ 338.0 [1705.7 - 462.7]	1412.6 $\pm$ 1123.8 [3421.6 - 457.5]	1057.9 $\pm$ 189.9 [1312.3 - 654.7]	704.5 $\pm$ 241.6 [1145.9 - 448.2]
C5	894.6 $\pm$ 390.4 [1530.1 - 226.8]	798.8 $\pm$ 488.5 [1530.1 - 226.8]	1053.7 $\pm$ 367.3 [1705.7 - 462.7]	785.5 $\pm$ 136.3 [970.3 - 649.1]
C6	1211.9 $\pm$ 635.5 [2665.0 - 457.5]	n.a	n.a	n.a

442



443

444 Figure 5: Polar Plots of BC concentration for the whole period (upper panel) and in three seasons (winter,  
 445 summer, and autumn) at Kenitra city. The average of BC concentrations ( $\text{ng}/\text{m}^3$ ) is coded as a colour  
 446 scale shown to the right. The dashed circular grey lines show the wind speed scale.

### 447 3.4 Diurnal BC concentrations patterns

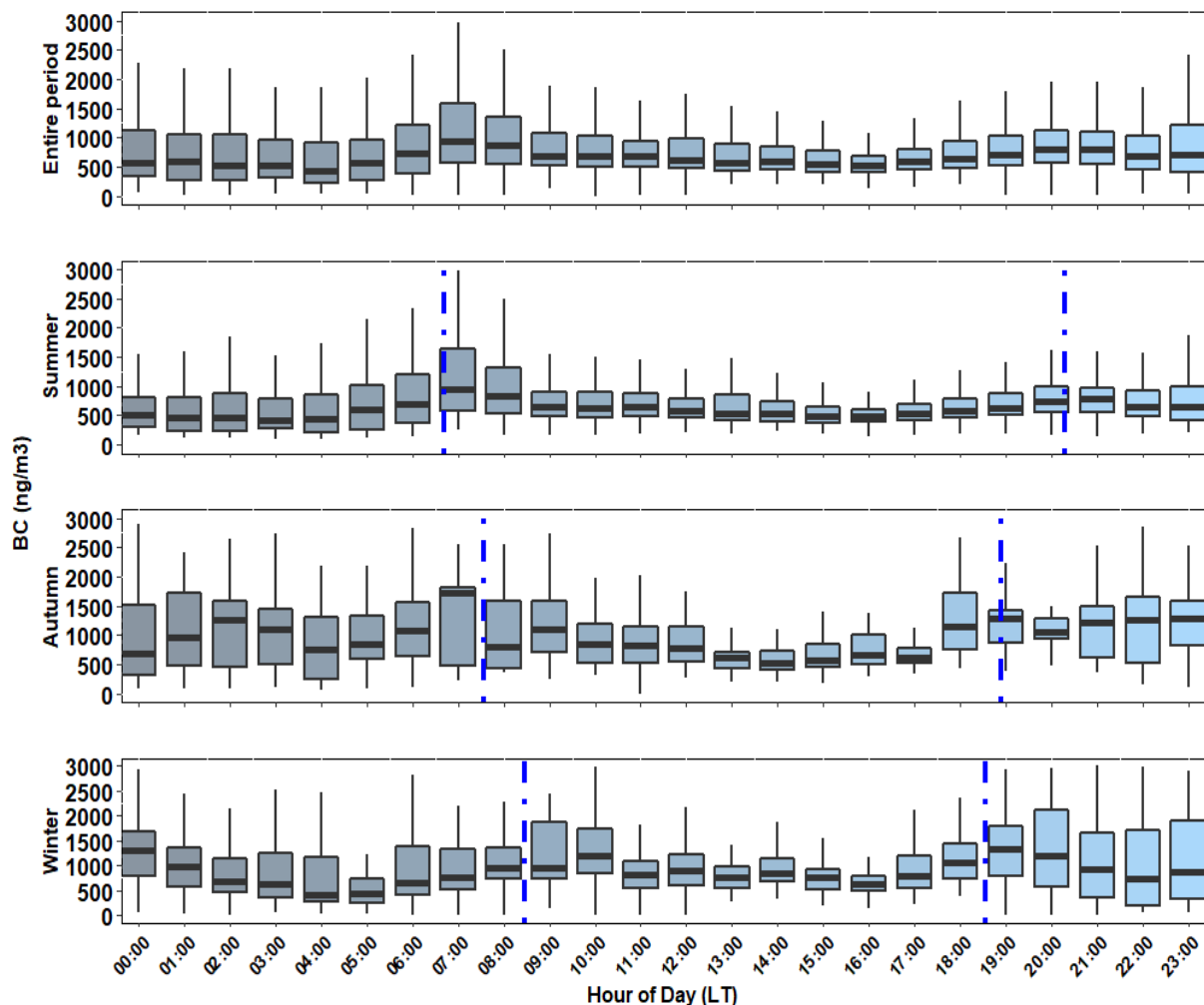
448 Besides the seasonal averages, the profile of BC concentration has also been evaluated on a higher  
 449 temporal resolution. Diurnal cycles of BC have been calculated as 1-hour mean values and are shown  
 450 in Figures 6a and 6b. From the temporal variation plots of hourly-averaged BC concentrations, a well-  
 451 defined and significant diurnal variation was observed throughout the day. Overall, a steady build-up in  
 452 the BC concentration to sharp peaks between 06:00 and 09:00 a.m local time (LT) just after sunrise  
 453 preceded by a smooth decrease to low concentration at midday (11:00 a.m - 05:00 p.m) and then again  
 454 starts increasing to attain another peak value at evening around 09:00 p.m after sunset. Table 2 shows  
 455 the sunrise and sunset time in different seasons. During the whole period, mean hourly BC  
 456 concentrations were found to range between 571.7 and 1311.6  $\text{ng}/\text{m}^3$  with an average value of 908.8  
 457  $\text{ng}/\text{m}^3$ . In summer, the average diurnal profile was fairly flat, with a sharp peak in the morning centered  
 458 on 06:00 to 09:00 a.m, accompanied by decreasing trend to minimum values reached at about 03:00 to  
 459 04:00 p.m, and then followed by gradual ascending with broader evening peak appeared at 09:00 p.m.  
 460 The BC level was still maintained until midnight. The average hourly BC concentration during summer  
 461 is 758.1  $\text{ng}/\text{m}^3$  and ranged from 516.9 to 1365.6  $\text{ng}/\text{m}^3$ . During winter, the average diurnal profile  
 462 showed a pre-dawn minimum at 5:00 a.m followed by a wide peak in the morning from 6:00 a.m to  
 463 12:00 p.m and a second peak, relatively more pronounced, during 06:00-11:00 p.m. The BC  
 464 concentration during the night-time followed an increasing pattern until 02:00 a.m, and the maximum  
 465 mean concentration exceeded nearly 2000  $\text{ng}/\text{m}^3$ . Maximum hourly BC concentrations in the morning  
 466 were about 1365.6 (at 07:00 a.m), 1432.9 (at 08:00 a.m), and 1548.3  $\text{ng}/\text{m}^3$  (at 10:00 a.m), and at evening  
 467 were around 952.1 (at 09:00 p.m), 1482.7 (at 11:00 p.m) and 2004.7  $\text{ng}/\text{m}^3$  (at 10:00 p.m) during  
 468 summer, autumn, and winter, respectively.

469 These indicate that the highest values were found during the winter and the lowest values were obtained  
470 during the summer, which is consistent with the seasonal changes previously discussed. The diurnal  
471 amplitudes (ratio of night-time peak to the daily minimum) of BC values were found to be seasonally  
472 dependent with the maximum occurring in winter (~3.1), followed by autumn (~2.4), and the lowest  
473 values observed during summer (~1.8). The mean annual variation amplitude of ~ 2.0 is of the order of  
474 magnitude observed in semi-urban Nagpur, Central India (Kompalli et al., 2014b). Hourly changes in  
475 BC concentrations throughout the autumn season indicated also bimodal distributions alike winter and  
476 summer seasons, while they were found to be the smallest during midday. The ratios of BC  
477 concentrations between winter and summer exhibit that the winter BC level is about a factor of 1.5 - 2.7  
478 higher than those measured in summer during the morning and of 2.2 - 2.9 during the late evening.  
479 During the afternoon hours, this ratio increases to about 2.8 to 4.0. The striking feature is that the average  
480 ratio between night-time (07:00 p.m to 06:00 a.m) and daytime (07:00 a.m to 06:00 p.m) of BC  
481 concentrations, estimated during the whole period, was found to be larger than the unity (1.19) as shown  
482 in Table 2. However, this feature is not pronounced during evening hours (06:00 - 08:00 p.m) when BC  
483 concentrations observed are lower than in the morning, excepting the winter season (ratio = 1.37). Tiwari  
484 et al. (2013) suggested that ratios larger than the unity are likely to indicate that the evening peaks result  
485 from changes in emission sources and variability in meteorological conditions. Doumbia et al. (2012)  
486 advocated that this may be attributed to the strong contribution of local sources mainly commuter traffic  
487 and to the fumigation effect in the boundary layer, which brings the evening aerosols near the surface  
488 from the nocturnal residual boundary layer shortly after the sunrise. Srivastava et al. (2019) stated that  
489 the entrainment of nocturnal BC residues begins in the early hours of the morning and leads to an  
490 increased BC concentration for a few hours. This phenomenon, commonly referred to as the fumigation  
491 effect, coupled with other anthropogenic emissions during early rush hours, is likely responsible for the  
492 morning BC peaks. Season-wise, its ratio was found to be 1.49, 1.32, and 1.02 during the winter, autumn,  
493 and summer seasons, respectively, exhibiting certain seasonal differences, which may be attributed to  
494 the differences in emission intensity and meteorological conditions of the sources associated with BC  
495 in each season. The most noticeable feature is that most of the time, the morning BC concentrations are  
496 higher than those observed during the evening hours. These BC diurnal profile peaks observed in this  
497 study were very comparable to that documented by others at different locations as in Shenzhen by Sun  
498 et al. (2020b), Karachi by Bibi et al. (2017), Dakar and Bamako by Doumbia et al. (2012), New Delhi  
499 by Tiwari et al. (2013), and Pune by Safai et al. (2007), i.e., distinct seasonal hourly BC trends can be  
500 noticed throughout the entire day (winter>autumn>summer). In a summary, it was observed that the  
501 diurnal profile of BC concentration follows bimodal behaviour in all the seasons but with varying  
502 degrees of the magnitude of the peaks. Indeed, the striking feature is that the morning peaks were more  
503 pronounced as compared with evening ones, and peaks were highest in winter and lowest in summer.

504 The fumigation effect leads to morning peaks in BC concentrations. The thermals generated after sunrise  
505 break the night-time inversion and bring down the BC from the residual layer to the surface, resulting  
506 in a peak. Concentration in rush hour traffic due to the build-up of vehicle traffic also contributes to this.  
507 Following this, enhanced convective mixing together with a deepening of the BLH leads to a drop in  
508 BC concentrations at the surface. As the thermals subside after sunset, the shallow nocturnal boundary  
509 layer sets in and the resulting stable conditions result in BC confinement near the surface. In addition,  
510 typically lower temperatures and wind speeds coupled with reduced source strength (less traffic during  
511 night-time curfews) lead to a gradual decrease in BC mass concentration leading to a night-time  
512 minimum. Seasonal differences in fumigation amplitudes and nocturnal BC peaks are driven by reduced  
513 source strengths and the evolution of BLH.

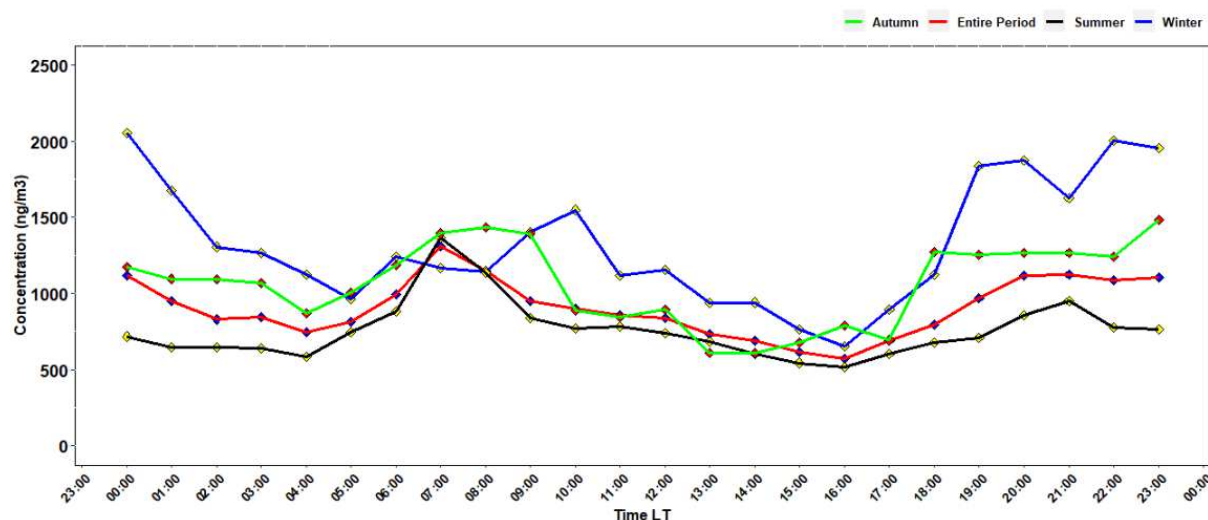
514 Over Kenitra, the BC concentration variations matched well with variations in diurnal anthropogenic  
515 activity emissions, population living style, and boundary layer development. Several studies (Liu et al.,  
516 2020; Tiwari et al., 2013; Doumbia et al., 2012; Safai et al., 2007) stated that the relationship between  
517 the BC concentrations and boundary layer height exhibits a negative correlation where the increased BC  
518 levels resulted in the lower boundary layer height. Ramachandran and Rajesh (2007) reported that the  
519 nocturnal boundary layer is three times narrower than that during daytime. Morning and evening BC  
520 peaks are likely driven primarily by anthropogenic activities (i.e. biomass burning, industrial activities,  
521 vehicular emission) and low boundary layer height during the morning and evening hours. Accordingly,

522 both increased local traffic emissions and the nocturnal boundary layer brake, which lifts aerosols to the  
 523 surface shortly after sunrise, contributed to the morning peaks. At noon hours, low BC concentrations  
 524 are attributed to i) aerosol dispersion due to the increase in surface temperatures which causes convective  
 525 activity (i.e., mixing) and elevation of the boundary layer (i.e., dilution), and ii) gradual decrease in  
 526 traffic volume after the morning rush hour. Later, the BC concentrations started building up gradually  
 527 and evening peaks started reappearing due to the commuter traffic activities (back home) that occurred  
 528 during this time and the decadence of the boundary layer. Ezani et al. (2021) argued that at sunset both  
 529 the boundary layer decreases and the atmospheric stability increases may lead to BC concentrations  
 530 increasing again. Then after, as shown in Figures 6a and 6b, the BC concentrations went on decreasing  
 531 slowly during the night-time but the BC levels were high which can be partly explained by trucks and  
 532 freight traffic either on the national way or the highway. As shown in Figure 1, the sampling site is  
 533 surrounded by three major traffic roads with distances between the site and these highways/corridors  
 534 ranging from 2.6 to 3.7 Km. As trucks typically have higher emission factors, which, when combined  
 535 with the general shadow boundary layer height during the evening leads to the enrichment of BC near  
 536 the surface (i.e. the evening peak) (Sun et al., 2020b). In brief, the observed diurnal variation of BC  
 537 concentrations is mainly attributed to the dynamics of the boundary layer height, although the urban and  
 538 local traffic activities might be contributing to the morning and nocturnal peaks.



539

540 Figure 6a: The diurnal distribution of BC during summer (mid-July - August), winter (December – mid-  
 541 February), and autumn (September - November) seasons, and the whole sampling period (upper panel).  
 542 The horizontal lines represent the medians; the limits of the boxes are the 1<sup>st</sup> and 3<sup>rd</sup> quartiles. The  
 543 whiskers extend to one and a half times the interquartile range. Dotted vertical lines represent the  
 544 seasonally mean sunrise and sunset times. Hours are in local time, LT.



545

546 Figure 6b: Hourly changes in BC concentrations throughout summer, winter, and autumn seasons.  
 547 Plotted points are the average values for aggregated data from the subsequent one hour on available days  
 548 (e.g., 10:00 LT represents the time from 10:00 LT to 10:59 LT). LT stands for local time.

549 **4. Future outlook**

550 This study aims at providing insights into an open gap in the scientific community, which is the spatial  
 551 variability of the black carbon aerosol in different environments and at different sites. We provide novel  
 552 data from North Africa (Morocco), with an emphasis on particular local conditions that multiply the BC  
 553 concentrations and consequently can affect human health. We finally participate in research efforts to  
 554 monitor carbonaceous aerosol emissions to inform the public of potential exposure.

555 This study is expected to provide valuable information to understand BC levels in North African urban  
 556 areas, relevant temporal variability, and emission sources and put this data into a larger perspective. It  
 557 also offers a reference for future attempts to assess fine carbonaceous aerosols in urban African areas.  
 558 Future outlook studies should concentrate on the role of local anthropogenic activities as compared to  
 559 long-range transport, and in particular on the impact of changes in regional characteristics on BC  
 560 concentrations in the Mediterranean areas; and BC source apportionment into wood burning (WB) and  
 561 fossil fuel combustion (FF) to investigate the quantitative contribution of such non-traffic sources to BC  
 562 (manuscript in preparation, comparing data between Kenitra, Morocco and Rome, Italy).

563 Limitations of this study include the sample size and non-uniform data spans, and these cannot be  
 564 changed due to variations in observational records. Future studies should add by performing longer (i.e.,  
 565 at least one-year) period characterization and trend analysis of BC concentrations, and using our data as  
 566 reference for previously available and valid observations.

567 **5. Conclusion**

568 The findings of this study provide the first attempt to assess long-term BC levels with state-of-the-art  
 569 technologies at a North-African urban area, Kenitra, in western Morocco. Data were collected during  
 570 mi-July 2020 - mi-February 2021. Based on average BC levels in the last 10 years, Kenitra represents a  
 571 low-to-medium BC concentration urban area for Africa, where data as high as more the 6 times were  
 572 found. Whether this is due to the peculiar COVID-19 period, different emission sources, or the different  
 573 measurement methods used will be the subject of future research.

574 In addition to these previous studies, the continuous dataset collected in this work with higher time  
 575 resolution (5-minutes continuous data for almost one year vs. 24-hours sporadic measurements) shed

576 light on diurnal, monthly, and seasonal variability in BC mass concentration over this African urban  
577 location (Kenitra).

578 The major findings obtained are summarized as follows:

- 579 - The average BC concentration observed in this study ( $0.90 \pm 0.80 \mu\text{g}/\text{m}^3$ ) is overall lower than  
580 those reported in African and European urban locations. This observed low carbonaceous  
581 pollution levels may be partially attributed to the reduction of anthropogenic activities due to  
582 the COVID19 curfew imposed by the authorities.
- 583 - BC concentrations displayed a clear bimodal diurnal pattern, in which morning peaks were  
584 attributed to morning inversion and rush-hour traffic, while evening peaks were related to  
585 combustion sources (evening traffic rush-hour) and the impacts of a shallow night-time  
586 boundary layer.
- 587 - The average BC values at night-time (07:00 p.m to 06:00 a.m) are 1.9 times greater than those  
588 observed throughout the day (07:00 a.m to 06:00 p.m) as the atmospheric boundary layer is  
589 shallower at night, retaining pollutants in a smaller volume, which leads to higher BC mass  
590 concentrations.
- 591 - BC exhibits a strong seasonal variability, ranging from a low of  $0.75 \mu\text{g}/\text{m}^3$  in the summer to a  
592 high of  $1.30 \mu\text{g}/\text{m}^3$  in the winter. Although lower BC concentration occurs in the summer due  
593 to aerosol dispersion driven by high wind speed, higher BC arises during winter due to shallow  
594 atmospheric boundary layer, long-range transport, and local emissions associated with low wind  
595 speeds.
- 596 - The association of BC with low wind speeds without marked directionality is a clear indication  
597 that the major source of BC is in the vicinity of the monitoring site, consistent with road traffic.  
598 Moreover, long-range westerly atmospheric transport events can contribute to elevated BC  
599 concentrations.
- 600 - During the afternoon period (11:00 a.m - 04:00 p.m), the BC concentration values are minimum  
601 (do not exceed the value of  $500 \text{ ng}/\text{m}^3$ , on average), because of the fully evolved boundary layer  
602 and reduced traffic activities, representing somehow an urban background level in  
603 Kenitra.

#### 604 **Author contributions**

605 Youssef Bounakhla: Performing the experiment, Writing - original draft preparation. Abdelfettah  
606 Benchrif: Conceptualization, Data Analysis, Writing - original draft, Supervision. Mounia Tahri:  
607 Writing - review & editing. Francesca Costabile: Conceptualization, Writing - review & editing. Fatiha  
608 Zahry: Performing the experiment. Moussa Bounakhla: Supervision. El Kafssaoui El Hassan:  
609 Supervision. All authors have read and agreed to the published version of the manuscript.

#### 610 **Competing interests**

611 The authors declare that they have no conflict of interest.

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